

1 **Contrasting physiological responses to future ocean acidification**
2 **among Arctic copepod populations**

3 Running head: Contrasting responses to ocean acidification

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29 Abstract

30 Widespread ocean acidification (OA) is modifying the chemistry of the global ocean, and the
31 Arctic is recognised as the region where the changes will progress at the fastest rate.
32 Moreover, Arctic species show lower capacity for cellular homeostasis and acid-base
33 regulation rendering them particularly vulnerable to OA. In the present study, we found
34 physiological differences in OA response across geographically separated populations of
35 *Calanus glacialis*. In copepodite stage CIV, measured reaction norms of ingestion rate and
36 metabolic rate showed severe reductions in ingestion and increased metabolic expenses in two
37 populations from Svalbard (Kongsfjord and Billefjord) whereas no effects were observed in a
38 population from the Disko Bay, West Greenland. At pH_T 7.87, which has been predicted for
39 the Svalbard west coast by year 2100, these changes resulted in reductions in scope for
40 growth of 19% in the Kongsfjord and a staggering 50% in the Billefjord. Interestingly, these
41 effects were not observed in stage CV copepodites from any of the three locations. It seems
42 that CVs may be more tolerant to OA perhaps due to a general physiological reorganisation to
43 meet low intracellular pH during hibernation. Needless to say, the observed changes in the
44 CIV stage will have serious implications for the *C. glacialis* population health status and
45 growth around Svalbard. However, OA tolerant populations such as the one in the Disko Bay
46 could help to alleviate severe effects in *C. glacialis* as a species.

47 Introduction

48 Widespread ocean acidification (OA) is modifying the chemistry of the global ocean (Hoegh-
49 Guldberg *et al.*, 2014). Driven by an increase in global atmospheric $p\text{CO}_2$ from 280 μatm at
50 pre-industrial times to the present day 400 μatm (IPCC, 2013), the global ocean mean surface
51 pH has decreased from 8.13 to the present day 8.05. Ocean models predict a continuation of
52 this trend with a further decrease of 0.4 pH units by the year 2100 (Bopp *et al.*, 2013, Caldeira
53 & Wickett, 2005, Cao *et al.*, 2007). Due to the chemical characteristics of Arctic sea water,
54 the Arctic is recognised as the region where the earliest and strongest decreases in pH are
55 expected (Fabry *et al.*, 2009, Hoegh-Guldberg *et al.*, 2014, Steinacher *et al.*, 2009). Increasing
56 sea ice melt with low H^+ buffering capacity makes Arctic waters increasingly susceptible to
57 OA (Yamamoto-Kawai *et al.*, 2009). Moreover, while the Arctic Ocean constitutes only 1%
58 of the global ocean volume, it receives 11% of the riverine discharge carrying not only low H^+
59 buffering capacity but also significant loads of terrestrial carbon prone to conversion to CO_2
60 by microbial respiration (Raymond *et al.*, 2007). This input has increased by 7% since the

61 1930s (Peterson *et al.*, 2002). Finally, increasing inflow from the North Atlantic carries large
62 amounts of anthropogenic CO₂ to the Arctic Ocean (Fransson *et al.*, 2001).

63 The magnitude of predicted chemical changes due to OA extends beyond anything
64 experienced by most extant species (Fabry *et al.*, 2008) and significant effects are predicted
65 for many marine animals (Dupont & Pörtner, 2013, Wittmann & Pörtner, 2013). But while
66 effects may be severe locally, they may vary across geographic ranges and among populations
67 (Wood *et al.*, 2016). While it has long been hypothesised that long distance dispersal of
68 planktonic larvae and eggs in an environment with few physical barriers has rendered most
69 marine species genetically homogeneous over long distances, recent studies of marine
70 invertebrates, including planktonic species, show geographically structured populations and
71 isolation on the scale of ocean basins and adjacent seas (Hellberg, 2009, Peijnenburg &
72 Goetze, 2013, Sanford & Kelly, 2010). Such structuring increases the possibility for
73 differential physiological responses to environmental changes to develop among hydrographic
74 provinces (as shown at lower latitudes by Calosi *et al.*, 2017, Vargas *et al.*, 2017). Differential
75 responses carry with them a possibility that affected species may be relieved from severe
76 effects and extinction (Calosi *et al.*, 2016, Sunday *et al.*, 2014). Effects may be severe locally,
77 and possibly lead to local extinction, but other enclaves may show higher tolerance.

78 Naturally, relief from environmental change is all the more important for the future of more
79 environmentally sensitive species, and energetic studies suggest that the capacity to counter
80 negative effects of OA could be particularly low in Arctic species. Contrary to cold adapted
81 eurythermal animals, true Polar species show low energetic costs for maintenance (Clarke,
82 1980, Rastrick & Whiteley, 2011). While this is an evolutionary strategy to enhance growth at
83 limited aerobic scope, lower allocation to cover maintenance costs also reduce the capacity
84 for energy demanding cellular homeostasis and acid-base regulation (Whiteley, 2011).
85 Moreover, because Arctic communities are characterised by simpler food webs – fewer
86 trophic levels and fewer species occupying each trophic level – they experience reduced
87 overall resilience to environmental changes (AMAP, 2013).

88 Calanoid copepods, particularly of the *Calanus* genus, constitute keystone species in the
89 Arctic pelagic community (Grainger, 1965, Møller *et al.*, 2006, Thor *et al.*, 2005). In most
90 pelagic communities, these crustaceans constitute 80% of the zooplankton biomass, and they
91 are the dominant component of prey for the larvae of most fish species (Last, 1980).
92 Consequently, their presence is fundamental to many fish populations and studies have shown
93 that larval survival and recruitment of such species as cod (*Gadus morhua*) and mackerel

94 (*Scomber scombrus*) co-vary with copepod abundance and biomass (Beaugrand *et al.*, 2003,
95 Castonguay *et al.*, 2008, Runge *et al.*, 1999). Any negative effects of environmental changes
96 will therefore have severe repercussions far beyond the copepod populations themselves. For
97 instance, increase in rainfall since the 1980s and lack of intrusion of high saline water from
98 the North Sea have affected reproduction and maturation in the copepod *Pseudocalanus*
99 *elongatus* in the Baltic Sea deep basins (Möllmann *et al.*, 2003). This has forced herring
100 (*Clupea harengus*) to revert to less favourable prey imposing serious implications for their
101 development and population growth (Möllmann *et al.*, 2003).

102 In the present study we investigated the possible existence of differential responses to OA
103 among geographically separated populations of *Calanus glacialis*, a species which dominates
104 the shelf of the Arctic Ocean and adjacent seas (Wassmann *et al.*, 2015). We established
105 physiological reaction norms across a pH gradient covering present and predicted future
106 environmental pH variability for Arctic continental shelf seas. Physiological response was
107 measured as the balance between energy intake and expenditure because it is this balance that
108 determines energetic performance and ultimately fitness in heterotrophs (Brown *et al.*, 2004).

109 **Methods**

110 **Collection of copepods**

111 Copepods were caught by vertical tows of a 200 µm WP2 net equipped with a closed cod end
112 from 100 m to the surface in the Kongsfjord, Svalbard (79.0° N, 11.7° E), the Billefjord,
113 Svalbard (78.6° N, 16.5° E), and the Disko Bay, Western Greenland (69°15' N, 53° 33' W)
114 during July 2015 (Fig. 1). On deck, the content of the cod end was diluted in 25 L seawater
115 collected at 80 m. Copepods were then transported to cold rooms (5 °C) at either the Kings
116 Bay Marine Laboratory (Ny-Ålesund, Svalbard) or the Arctic Station Laboratory
117 (Qeqertarsuaq, Western Greenland). *Calanus glacialis* copepodites stages III, IV, and V
118 (hereafter CIII, CIV, and CV) were selected under the stereomicroscope using cut off plastic
119 Pasteur pipettes, keeping all vessels on ice to avoid high temperatures. Copepodite stages
120 were identified by number of pleopods and abdominal segments (Mauchline, 1998). They
121 were distinguished from *Calanus hyperboreus* and *Calanus finmarchicus* copepodites on the
122 basis of prosome size (Arnkværn *et al.*, 2005, Thor *et al.*, 2008), by red pigmentation in the
123 antennules, which *C. finmarchicus* most often do not have (Nielsen *et al.*, 2014), and the lack
124 of lateral spikes on the distal prosome segment, which is a characteristic of *C. hyperboreus*
125 (Klekowski & Weslawski, 1991).

126 Experimental design

127 We applied a regression design approach, exposing independent samples of copepods to one
128 of seven to nine pH levels (Table 1). This approach has the advantage of enhanced predictive
129 power compared to the character state approach, which compares effects among different
130 distinct future climate scenarios (Havenhand *et al.*, 2010). We found CIIs only in the
131 Kongsfjord population, whilst CIVs and CVs were found at all three locations. However, CVs
132 were found in very low numbers in the Billefjord population. After removal of replicates
133 containing incorrectly stage determined individuals (as determined from photographs),
134 individuals with very aberrant prosome length also indicative of erroneous stage
135 determination or speciation, and individuals judged dead after incubations, a total of 153
136 replicates of ingestion rate measurements and 170 replicates of metabolic rate measurements
137 remained (Table 1).

138 Preparation of incubation water

139 For the initiation of incubations and at each water change, five litre batches of incubation
140 water for each treatment were prepared by mixing 0.3 μm filtered seawater (*fsw*) with small
141 volumes of *fsw* acidified to ca. pH 5.5 by CO₂ bubbling (Mapcon© CO₂, Yara Praxair,
142 Tromsø, Norway). This method for manipulating seawater carbonate chemistry has been
143 previously described and validated (Riebesell *et al.*, 2010). The different treatments were
144 established at target pH_T (pH on the total scale) increments of 0.2. Total alkalinity (A_T) was
145 analysed by potentiometric titration (Dickson *et al.*, 2007) in an open cell with 0.1 M HCl
146 using a VINDTA 042 carbonate titrator (Marianda, Germany) and total dissolved inorganic
147 carbon (C_T) was analysed by coulometric titration (Dickson *et al.*, 2007) using a coulometer
148 (CM5015, UIC, Joliet, IL, USA) connected to the VINDTA after acidification with 8.5 %
149 phosphoric acid. pCO₂ and pH_T were calculated using CO2SYS (Pierrot *et al.*, 2006) with
150 constants from Mehrbach *et al.* (1973) refitted by Dickson and Millero (1987) and inputs of
151 temperature, salinity, A_T, and C_T. pH_T was monitored using a SevenGo SG2 pH meter
152 equipped with an InLab 413 SG/2m electrode (Mettler-Toledo, Columbus, Ohio, USA)
153 (Svalbard populations) or a HI 98183 pH/ORP meter (Hanna, Woonsocket, Rhode Island,
154 USA) (Disko Bay population). Determination of pH_T in all incubation water batches and
155 incubation bottles were based on a standard curve established from simultaneous
156 measurements in water samples of electric potential (mV) with the pH electrodes and
157 determination of pH_T from A_T and C_T with the VINDTA in the pH range 8.2-6.4. Salinity and
158 temperature were measured using a conductimeter (Cond 340i, WTW, Weilheim, Germany).

159 Measured values of chemistry parameters are shown in Table 2. A_T was established only once
160 for the Billefjord population. For food, paste of the diatom *Thalassiosira weissflogii* (Tw
161 1200, Reed Mariculture, Campbell, CA, USA) was added to a final concentration of ca. 10 μg
162 Chl *a* L^{-1} . The necessary dilution of the algal paste was established from the Chl *a* content of
163 the algal paste determined spectrophotometrically (UV-2401 PC, Shimadzu Co., Kyoto,
164 Japan) after overnight extraction in 70% ethanol (Strickland & Parsons, 1972). Prior to
165 incubations, the suitability of the algal paste as prey for *C. glacialis* was assured by
166 comparing faecal pellet counts from incubations of copepodites with previous counts from
167 copepodites incubated at similar concentrations of algae.

168 Copepod incubations

169 For each experiment, copepodites were incubated for a total of 8 d (7 d incubation plus 1 day
170 ingestion rate measurements). For each replicate, 10 individuals were pipetted, using cut off
171 plastic Pasteur pipettes, into a 600 mL glass Duran bottles prepared with incubation water. All
172 bottles were closed, making sure no air bubbles were present, and placed on a slowly rotating
173 plankton wheel (0.5 rpm) at ca. 5 °C in dim light. Every day approximately 500 mL water was
174 replaced in each bottle by inserting a piece of pipe fitted with a 200 μm screen at the bottom,
175 siphoning off the water from inside the tube, and replacing it with water from the pre-
176 prepared five litre incubation water batches at the appropriate pH. Samples for A_T and C_T
177 were taken from the incubation water batches and from water pooled from all bottles of each
178 treatment subsequent to the incubations on days 2, 5, and 8).

179 Measurement of ingestion and metabolic rates

180 On day 7, five additional control bottles without copepods were prepared with incubation
181 water for estimates of ingestion rates. Triplicate samples for Chl *a* determination were taken
182 from each incubation water batch. On day 8 the content of each bottle was poured through a
183 20 μm sieve held in a Petri dish to remove copepods, faecal pellets, and eggs. While doing
184 this, the water was collected in a beaker from under the Petri dish and 200 mL was filtered
185 onto a 0.7 μm glass fiber filter (Whatman, GF/F, Maidstone, UK) which was frozen for later
186 Chl *a* determination. The content of the 20 μm sieve was gently flushed into a Petri dish and
187 copepods for metabolic rate measurements were collected. The rest were counted and
188 photographed for precise determination of developmental stage under the stereoscope.

189 For estimates of specific metabolic rate ($\dot{M}O_2$), oxygen consumption rates were measured on
190 individual copepodites according to Thor and Oliva (2015). One individual from each bottle

191 was pipetted from the Petri dish into a 1.6 mL vial fitted with fluorescent O₂ reactive foil
192 discs (PSt3 spots, PreSens, Regensburg, Germany) and filled with *fsw*, which had been
193 saturated with air by vigorous bubbling and adjusted to the corresponding pH. Vials were then
194 sealed with Teflon caps and after a resting period of ca. 30 min to acclimate copepods O₂
195 concentrations were measured at 0, 2.5, and 5 h using an optode O₂ system (Fibox 3, PreSens,
196 Regensburg, Germany). O₂ consumption rate (nmol O₂ ind⁻¹ d⁻¹) was calculated by subtracting
197 the average O₂ depletion rate measured in the five controls without copepods from the O₂
198 depletion rate in each of the copepod containing vials (nmol O₂ L⁻¹ h⁻¹) and multiplying by
199 vial volume (L) and 24 h d⁻¹. Prior testing of the optode system at 5 °C showed a 3-min 95 %
200 reaction time, i.e. the period of time taken before the output reached within 5 % of the final O₂
201 concentration value (as estimated by exponential regression). Therefore, at every sampling
202 event, O₂ concentration was read for 3 min, and an average of values read during the last
203 minute was used for calculations. Subsequent to the measurements the copepods were
204 transferred to Petri dishes and photographed under the stereoscope for detailed stage
205 determination.

206 For estimates of ingestion rate, phytoplankton Chl *a* concentrations of all samples were
207 determined fluorometrically. The frozen filters were extracted in 4 mL acetone overnight and
208 fluorescence was measured on a Turner Designs 10-AU fluorometer (Strickland & Parsons,
209 1972). Ingestion rate (µg Chl *a* ind⁻¹ d⁻¹) was calculated from the decrease in Chl *a*
210 concentrations from all bottles containing copepods subtracted by the decrease in
211 disappearance from the control bottles (µg Chl *a* L⁻¹ d⁻¹) (Frost, 1972), multiplying by bottle
212 volume (L), and dividing by number of copepods counted in the bottles at day 8.

213 To obtain weight specific rates, copepod prosome lengths were measured from the
214 photographs using ImageJ (U. S. National Institutes of Health) and body carbon weights were
215 calculated using a weight/length relationship of $W (\mu\text{gC}) = 4.8L (\text{mm})^{3.57}$ (Madsen *et al.*,
216 2001). Oxygen consumption rates (nmol O₂ ind⁻¹ h⁻¹) were converted to specific metabolic
217 rate ($\dot{M}O_2$, µgC µgC⁻¹ d⁻¹) by dividing by body mass (µgC ind⁻¹), multiplying by a respiratory
218 coefficient of 0.97 mol C mol O₂⁻¹ (Omori & Ikeda, 1984), multiplying by 0.012 µgC nmol C⁻¹
219 ¹, and multiplying by 24 h d⁻¹. Ingestion rates (ng Chl *a* ind⁻¹ d⁻¹) were converted to specific
220 ingestion rate (*IR*, µgC µgC⁻¹ d⁻¹) by multiplying by 50 µgC µg Chl *a*⁻¹ (Båmstedt *et al.*,
221 2000) and dividing by body mass (µgC ind⁻¹).

222 To avoid bias from differences in temperature among incubations, all rates were normalized
223 to the average temperature of 5.2 °C using a Q_{10} value of 2.0 for metabolic rate in marine
224 copepods (Ikeda *et al.*, 2001).

225 Data analysis and determination of reaction norms

226 Since treatments were evenly distributed along pH reaction norms for each population and
227 copepodite stage, rates would be inherently non-normally distributed when reaction norms
228 show significant slopes. For comparisons of mean rates (i.e. the average rate of all individuals
229 from all pH treatments) among populations and stages we therefore used a 2-factor
230 permutational analysis of variance test (PERMANOVA) on similarity matrices assembled
231 using Euclidian distances (Anderson, 2001). Prosome lengths were similarly compared among
232 populations and stages using a 2-factor PERMANOVA.

233 For each copepodite stage in each population, pH reaction norms of ingestion rate and
234 metabolic rate were established by sequentially testing polynomial regression models of
235 increasing order (linear, quadratic, or cubic) for the relationship between the variable and pH_T
236 according to David *et al.* (1997). Best fitting models were chosen by statistically comparing
237 sums of squares among the three models as

$$238 \quad F_{1,df} = \frac{SS_{higher} - SS_{lower}}{MS_{res}}$$

239 where df is the degree of freedom of the higher degree model, SS_{higher} is the sums of squares
240 of the higher degree model, SS_{lower} is the sums of squares of the lower degree model, and
241 MS_{res} is the residual mean squares of the higher degree model (Rocha & Klaczko, 2012).

242 After assuring homoscedasticity (Levene's test), reaction norms of specific rates were
243 compared among populations using univariate general linear model analysis (GLM) in SPSS
244 (IBM Inc.). Differences of level among populations were detected by significant differences
245 among populations using a pH_T + population design, and differences of slopes were detected
246 by significant interactions of pH_T and population using pH as the covariate in a pH_T +
247 population + population x pH_T design.

248 To evaluate the overall physiological effects of decreasing pH_T , scope for growth values were
249 constructed from relationships between metabolic rate and ingestion rate in CIVs. Since
250 metabolic rates were measured on different individuals than ingestion rate, no direct
251 comparison was possible and we therefore calculated mean predicted scope for growth values

252 (\widehat{SFG}) at each pH_T on the basis of predicted rates from the reaction norm regressions as
253 $\widehat{SFG} = \widehat{IR} \times AE - \widehat{MO}_2$, where AE is absorption efficiency, which was set at 0.6 for
254 copepods (Thor *et al.*, 2007, Thor & Wendt, 2010).

255 Results

256 Comparison of mean rates among populations and developmental stages

257 Although prosome lengths were measure purely to enable calculation of weight specific rates,
258 we found significant differences in these among populations (unrelated to pH) and therefore
259 report the analyses here. Prosome lengths of both stage CIV and CV copepodites differed
260 significantly among the three populations (2-factor PERMANOVA: pseudo- $F_{2,335} = 32.2$, $P <$
261 0.001). CIVs were significantly larger in the Kongsfjord and Disko Bay populations ($2532 \pm$
262 $381 \mu\text{m}$ and $2510 \pm 115 \mu\text{m}$, mean \pm sd), respectively, than in the Billefjord population (2338
263 $\pm 150 \mu\text{m}$) (2-factor PERMANOVA pair-wise test: $P < 0.001$), whereas CVs were
264 significantly larger in the Disko Bay population ($3357 \pm 144 \mu\text{m}$) than in the Kongsfjord and
265 Billefjord populations ($2962 \pm 307 \mu\text{m}$ and $2875 \pm 313 \mu\text{m}$, respectively) (2-factor
266 PERMANOVA pair-wise test, $P < 0.001$).

267 The mean specific ingestion rate of the three developmental stages (for each stage, the
268 average rate of all individuals from all pH_T tested) were significantly different at $0.111 \pm$
269 $0.042 \mu\text{gC } \mu\text{gC}^{-1} \text{ d}^{-1}$ in CIIIs, $0.044 \pm 0.021 \mu\text{gC } \mu\text{gC}^{-1} \text{ d}^{-1}$ in CIVs, and $0.021 \pm 0.011 \mu\text{gC}$
270 $\mu\text{gC}^{-1} \text{ d}^{-1}$ in CVs (2-factor PERMANOVA: pseudo- $F_{2,152} = 54.6$, $P < 0.001$). Mean rates (for
271 each population, the average rate of all individuals from all pH_T tested) also differed
272 significantly between the Kongsfjord and Disko Bay populations (2-factor PERMANOVA
273 pairwise test: $P = 0.004$) mainly due to the larger size and calculated weight, and hence lower
274 specific rates, of CVs in the Disko Bay population.

275 Similarly, mean specific metabolic rates were significantly different among developmental
276 stages: $0.025 \pm 0.018 \mu\text{gC } \mu\text{gC}^{-1} \text{ d}^{-1}$ in CIIIs, $0.024 \pm 0.009 \mu\text{gC } \mu\text{gC}^{-1} \text{ d}^{-1}$ in CIVs, and 0.015
277 $\pm 0.006 \mu\text{gC } \mu\text{gC}^{-1} \text{ d}^{-1}$ in CVs (2-factor PERMANOVA: pseudo- $F_{2,169} = 14.3$, $P < 0.001$).
278 These differed among populations with significantly lower rates in the Disko Bay population
279 than in the two Svalbard populations (2-factor PERMANOVA pairwise tests: $P < 0.02$).

280 Ingestion rate reaction norms

281 In CIVs ingestion rates decreased by 85% and 66% from the highest to the lowest pH_T , in the
282 Kongsfjord and Billefjord populations respectively, but remained unchanged in CIV from the

283 Disko Bay population (Figs. 2a,b,c). Ingestion rate reaction norms showed linearly decreasing
284 rates with decreasing pH_T in CIVs from the Kongsfjord and Billefjord populations (Table 3).
285 There was no difference in slopes between the Kongsfjord and Billefjord populations (GLM,
286 comparison of slopes: $F_{1,52} = 0.61$, $P = 0.439$).

287 In CIIs from the Kongsfjord population, ingestion rates first increased by 53% from the
288 highest pH_T to pH_T 7.337 and then decreased to 33% at the lowest pH_T compared to the rate at
289 the highest pH_T (Fig. 2d). These changes were better fitted with the second order regression,
290 $IR = maxIR + g_2(pH_T - pH_{TmaxIR})^2$, where maximum ingestion rate ($maxIR$) was 0.124
291 $\mu gC \mu gC^{-1} d^{-1}$, pH_T at maximum ingestion rate (pH_{TmaxIR}) was 7.41, and the slope, g_2 , was -
292 0.099 ($r^2 = 0.39$, $P = 0.019$) (Fig. 2d).

293 There were no significant effect of pH_T on ingestion rates of CVs from any of the three
294 populations (Table 3; Fig 3).

295 Metabolic rate reaction norms

296 Metabolic rates increased by 136% and 127% from high to low pH_T in CIVs from the
297 Kongsfjord and Billefjord populations, respectively, but remained unchanged in CIVs from
298 the Disko Bay population (Figs. 2a,b,c). The metabolic reaction norms showed significant
299 linearly increasing metabolic rates in Kongsfjord and Billefjord CIVs (Table 4) but there were
300 no differences in slopes of metabolic rate reaction norms between in the Kongsfjord and
301 Billefjord population CIVs (GLM pairwise comparison of slopes: $F_{1,48} = 1.30$, $P = 0.260$),
302 Metabolic rates remained unchanged with decreasing pH_T in CIIs (Table 4; Fig. 2d), and in
303 CVs from all three populations (Table 4; Fig. 3).

304 Temperatures were generally lower in the Disko Bay experiments. Correction for temperature
305 differences among locations changed rates by an average 8 %. These corrections did not
306 significantly affect reaction norm slopes (GLM analysis comparing slopes of all reaction
307 norms with and without temperature corrections: $P < 0.05$).

308 Scope for growth

309 In CIVs, \widehat{SFG} decreased from 0.032 $\mu gC \mu gC^{-1} d^{-1}$ at pH_T 8.012 to -0.021 $\mu gC \mu gC^{-1} d^{-1}$ at
310 pH_T 6.445 in the Kongsfjord population and from 0.010 at pH_T 8.041 to -0.018 $\mu gC \mu gC^{-1} d^{-1}$
311 at pH_T 7.036 in the Billefjord population. Thus, \widehat{SFG} became negative below pH_T 7.04 in
312 CIVs from the Kongsfjord population but already at pH_T 7.67 in CIVs from the Billefjord
313 population.

314 In CIIIs from the Kongsfjord population predicted scope for growth (\widehat{SFG}) first increased
315 from $0.025 \mu\text{gC } \mu\text{gC}^{-1} \text{ d}^{-1}$ at pH_T 8.041 to $0.049 \mu\text{gC } \mu\text{gC}^{-1} \text{ d}^{-1}$ at pH_T 7.333 and then
316 decreased to $-0.009 \mu\text{gC } \mu\text{gC}^{-1} \text{ d}^{-1}$ at pH_T 6.421.

317 We did not calculate \widehat{SFG} changes in CVs since neither ingestion rates nor metabolic rates
318 changed significantly with pH_T . Any calculated differences would stem from stochastic
319 differences or sampling variability rather than real physiological changes.

320 Discussion

321 The balance between energy intake and energy expenditure is the prime determinant of
322 survival in any heterotrophic organism. Energy intake has to be sufficient to cover
323 maintenance and repair costs, as well as costs for growth or reproduction for an organism to
324 uphold positive Darwinian fitness (Sibly & Calow, 1986). In the present study, we observed
325 severe reductions in ingestion rate along with increased metabolic rates with decreasing pH_T
326 in *Calanus glacialis* copepodite stage CIV from two Svalbard populations (Kongsfjord and
327 Billefjord), but not in CIVs from the Disko Bay, West Greenland. These effects were limited
328 to the CIV stage and there were no effects in stage CV copepodites from any of the three
329 populations. Nevertheless, at pH_T 7.87, which has been predicted for the Svalbard west coast
330 by the year 2100 (Bellerby *et al.*, 2012), scope for growth decreased by 19% in the
331 Kongsfjord CIVs, while in the Billefjord CIVs it decreased by a staggering 50%. In fact, these
332 estimates of scope for growth may be conservative since absorption efficiency may decrease
333 with decreasing pH due to decreasing gut enzyme activity (Stumpp *et al.*, 2013). Needless to
334 say, such changes will have serious implications for the *C. glacialis* population around
335 Svalbard. Reductions in scope for growth on this scale will prolong stage development time
336 and reduce the individual body size of the developing copepodites and ultimately also reduce
337 adult body size. This effect has been observed in *Calanus helgolandicus* cohorts reared in
338 mesocosms at low prey levels (Rey-Rassat *et al.*, 2002). The resulting reduction in adult body
339 size will entail decreased egg production rates (Halvorsen, 2015), and there is a real risk that
340 these effects, although possibly limited to one or a few specific copepodite stages (Kongsfjord
341 CIIIs showed a peaking ingestion rate reaction norm), may impair the general health status
342 and growth of *C. glacialis* in this region. Accordingly, studies in the North Sea and the sub-
343 Arctic Pacific have shown that similar changes in spring juvenile production have significant
344 effects on overall population development. A long-term sampling series in the North Sea has
345 shown that years with low larval growth during spring results in lower summer biomass than

346 years with higher spring larval growth (Clark *et al.*, 2003). Similar variations have been
347 observed in the sub-Arctic Pacific *Neocalanus plumchrus* population. This population
348 experiences significant inter-decadal variations in peak summer biomass, which is
349 hypothesised to stem from changes in copepodite growth rate during spring (Mackas *et al.*,
350 1998).

351 Previous studies have shown metabolic effects of low pH on copepods, although results are
352 far from conclusive. Metabolic rate increased significantly by 28% from pH_{NBS} (National
353 Bureau of Standards scale) 8.18 to 7.83 in *Centropages tenuiremis* (no developmental stage
354 indicated) (Li & Gao, 2012) and in *Pseudocalanus acuspes* females it increased significantly
355 by 11% from pH_T 8.06 to 7.75 (Thor & Dupont, 2015). Metabolic rates doubled from pH_T
356 8.06 to pH_T 7.66 in *Acartia grani* females, although low replication rendered the difference
357 non-significant, whereas no clear effect was observed in female *A. clausi* exposed to pH_T 8.03
358 and pH_T 7.83 (Isari *et al.*, 2015, Zervoudaki *et al.*, 2014). In *Pseudocalanus acuspes* a
359 decrease from 7.95 pH_T to 7.61 showed no clear effect on metabolic rate in a population from
360 Svalbard, whereas a population from Skagerrak experienced significant changes (Thor &
361 Oliva, 2015). But these changes depended on food level and no clear response could be
362 concluded. The lack of response of *C. glacialis* CVs in the present study is corroborated by a
363 recent study in the Kongsfjord (Thor *et al.*, 2016) and has also been shown to last during
364 longer-term incubations where metabolic rates remained equal in *C. glacialis* CVs and *C.*
365 *hyperboreus* CVs and females incubated at pH_F (free scale pH) 8.13 and 7.26 for 62 days
366 (Hildebrandt *et al.*, 2014). Metabolic rates of CVs increased linearly across a range from pH_T
367 8.02 to pH_T 7.16 in a study on culture reared *C. finmarchicus* applying reaction norm statistics
368 similar to the present study (Pedersen *et al.*, 2014), whereas a later study found no effects
369 between pH_T 7.92 and pH_T 7.51 in wild caught *C. finmarchicus* CVs and females (Runge *et*
370 *al.*, 2016). Ingestion rates have been shown to be unresponsive in *A. grani* and *Oithona*
371 *davisae* females (Isari *et al.*, 2015). In the *Calanus* genus, *C. finmarchicus* and *C. glacialis*
372 CVs showed no changes in ingestion rates when exposed at pH_T 7.2 (Hildebrandt *et al.*,
373 2016).

374 Geographically specific responses to low pH exposure have been demonstrated in several
375 marine species. The metabolic response to low pH varies with latitude in the gastropod
376 *Littorina littorea* showing an upregulation in the centre of the species distribution along the
377 European continental coast but a decrease in the southern- and northern-most regions (Calosi
378 *et al.*, 2017). Such latitudinal differences also occur in the calanoid copepod *Acartia tonsa*,

379 larvae of the gastropod *Concholepas concholepas*, and the bivalve *Perumytilus purpuratus*
380 along the Chilean coast (Vargas *et al.*, 2017). While ingestion rates did not change with
381 decreased pH in *A. tonsa* originating from an estuary with low and variable pH, they
382 decreased by 72% in individuals from a coastal ocean area with perpetual high pH (Vargas *et*
383 *al.*, 2017). Geographically specific responses have been observed also in another calanoid
384 copepod species, *Pseudocalanus acuspes*. Populations from the Kongsfjord and the
385 Gullmarsfjord (Swedish west coast) showed differences in the relationship between ingestion
386 rate and metabolic rate (Thor & Oliva, 2015). Low pH induced a steeper increase in metabolic
387 rate with increasing ingestion rate in females of the Swedish population than in females of the
388 Svalbard population. Also the isopod *Idotea balthica* has shown geographically specific OA
389 responses. In this case, metabolic rate and osmoregulatory activity responded differently to
390 increased $p\text{CO}_2$ (1000 μatm) in individuals originating from low and high salinity
391 environments (Wood *et al.*, 2016). Likewise, larvae of the spider crab *Hyas araneus* have
392 shown differences in growth responses between two populations from Svalbard and the North
393 Sea (Walther *et al.*, 2010). These differences may be a reflection of a general ability of the
394 tested species for physiological plasticity to counter pH variations. Such plasticity may
395 originate from the environment of the individual's habitat (phenotypic plasticity) or from the
396 environment experienced by previous generations (transgenerational plasticity). But they may
397 also arise from genetic adaptation to different pH environments among locations. Evidence
398 for rapid evolution in the face of fast environmental changes is increasing (Carroll *et al.*,
399 2007), and previous studies have shown that calanoid copepods have the capacity for fast
400 adaptation to low pH conditions. While our experimental design, incubations for less than one
401 generation, did not allow detection of local adaptation, Thor and Dupont (2015) found
402 adaptation causing changes in *Pseudocalanus acuspes* fecundity after only two generations at
403 pH_T 7.54, which could be linked to observed selection in genes coding for processes involved
404 in oxidative phosphorylation and ribosomal structure (De Wit *et al.*, 2015). Similarly, in
405 echinoderms low pH/high $p\text{CO}_2$ has been observed to induce rapid selection in genes coding
406 for biomineralization, lipid metabolism, and ion homeostasis (Pespeni *et al.*, 2013). However,
407 in the very same study on *P. acuspes*, Thor and Dupont (2015) also found evidence of
408 phenotypic plasticity in response to lowered pH, albeit at lower levels of pH reductions, so
409 both mechanisms may act in concert to alleviate OA effects. Regardless of the origin of the
410 observed geographic differences in the CIV copepodites, phenotypic plasticity,
411 transgenerational plasticity, or local adaptation, they have specific consequences for the future
412 of *C. glacialis* as a species. The severe reductions in scope for growth in this stage observed

413 in the Svalbard populations would render *C. glacialis* with little potential to survive future
414 OA. However, the existence of enclaves or perhaps extended populations with increased
415 tolerance, such as the Disko Bay population, could prove important as an alleviating factor to
416 remove or at least delay future OA effects.

417 Tolerance to certain environmental conditions is developed through pre-exposure. The few
418 existing studies reveal a possible difference between the Disko Bay and the Svalbard fjords
419 with respect to carbonate chemistry. While the Davis Strait outside Disko Bay exhibits similar
420 high pH, as is common in Arctic waters (Azetsu-Scott *et al.*, 2010), the water of the Disko
421 Bay may be somewhat special. The Disko Bay is influenced by extensive glacial discharge
422 from the Jakobshavn glacier, and during summer the surface water are characterised by the
423 balance between melt water production and the inflow of water from the West Greenland
424 Current (Hansen *et al.*, 2012). Hence, the Disko Bay is very variable environment both on a
425 seasonal and inter-annual scale. Studies from 2011 and 2012 showed that while pH_{NBS} was
426 mostly high at the surface, it was perpetually lower than 8.0 below 50 m with values
427 approaching 7.5 during May (Riisgaard *et al.*, 2015, Thoisen *et al.*, 2015). Frequently, low pH
428 water was encountered throughout the water column during May in both years studied. pH_{NBS}
429 did increase during the spring bloom but re-attained values below 8.0 immediately after the
430 termination of the bloom (Riisgaard *et al.*, 2015). Outside the spring bloom period, pH_{NBS} was
431 in the range 7.6-7.9 at fluorescence max depth, the depth where most copepods reside when
432 feeding. The Kongsfjord is probably the best studied of the three, and recent investigations
433 show high pH/low $p\text{CO}_2$ conditions throughout the fjord during summer and possibly also
434 during winter (Fransson *et al.*, 2016). pH_{T} remained above 8.0 throughout the water column
435 during July of the two consecutive years 2013 and 2014, and although winter data are scarcer,
436 minimum measured winter surface water pH_{T} values in the Kongsfjord were 8.11 in 2013 and
437 8.14 in 2014 (Fransson *et al.*, 2016). To our knowledge there is no information on carbonate
438 chemistry from the Billefjord. Thus, contrary to the Kongsfjord (and perhaps also the
439 Billefjord), it seems that there would be a real possibility for zooplankton in the Disko Bay to
440 be frequently exposed to low pH conditions during spring and summer, the period for
441 copepodite growth (Yamamoto-Kawai *et al.*, 2009).

442 Is tolerance of low pH a special characteristic of the Disko Bay population or could we expect
443 enclaves with similar tolerance elsewhere? While Arctic waters most often are characterised
444 by high pH, studies show that low pH conditions do develop temporarily in some areas.
445 Corrosive conditions have been observed in the Canada Basin connected to sea ice melt

446 (Yamamoto-Kawai *et al.*, 2009), and low pH/high pCO₂ conditions have also been observed
447 in extended areas along the Siberian coast (Anderson *et al.*, 2011). Here, in the Laptev Sea,
448 CO₂ produced from microbial decomposition of organic matter originating from river run-off
449 has been shown to oversaturate the entire water column, even in the post spring bloom period
450 (Anderson *et al.*, 2011). High pCO₂/low pH conditions have also been observed north of
451 Greenland (Jutterström & Anderson, 2010). Thus, these areas could potentially function to
452 pre-condition copepods to low or at least variable pH increasing the possibility of species
453 wide tolerance to future OA.

454 Because we studied different developmental stages, our findings also contributed another
455 important observation. While CIVs responded significantly to decreasing pH, we observed no
456 clear change in either ingestion or metabolic rate in CVs. Also in a previous study, Thor *et al.*
457 observed significant changes in the metabolic reaction to feeding at pH_T 7.73 compared to
458 pH_T 8.11 in early copepodite stages (CII-CIII) but no changes in CVs (Thor *et al.*, 2016).
459 Hildebrandt and colleagues found a similar lack of response of ingestion and metabolism in *C.*
460 *glacialis* CVs (Hildebrandt *et al.*, 2014, Hildebrandt *et al.*, 2016). But while this led the
461 authors to boldly conclude that shifts in seawater pH do not affect *C. glacialis* as a species,
462 our study highlights the need to refrain from conclusions based on studies of single
463 developmental stages. Such notion has been put forward previously by Dupont and colleagues
464 (2010). Their meta-analysis of OA effects in echinoderms showed that larvae and juveniles
465 mostly experience negative effects on growth and calcification while adults respond
466 positively. In crustaceans, stage-specific metabolic responses to OA were also found for
467 different larval stages in the European lobster (Small *et al.*, 2015). Also *Calanus* exhibits
468 fundamental stage-specific metabolic differences, and in this respect the CV stage stands out.
469 While somatic growth is the main goal in the preceding stages, metabolism is largely
470 reconfigured to accommodate overwintering diapause in CVs. Ingestion rates were not much
471 higher than metabolic expenses in this stage (Fig. 3) and it seems that CVs were entering this
472 phase of physiological reconfiguration at the time of measurements. During diapause, *C.*
473 *glacialis* CV experience extracellular pH as low as 5.5 possibly as a result of metabolic
474 depression during hibernation (Freese *et al.*, 2015). It is therefore quite conceivable that
475 mechanisms to counter low pH could be activated in this particular stage as part of the general
476 physiological reconfiguration to accommodate hibernation. This would render CVs
477 particularly unresponsive to ambient pH. If such mechanisms require energy, as most

478 physiological processes do, it would be evolutionarily beneficial to avoid their activation
479 before they are needed.

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493 References

- 494 Amap (2013) *AMAP Assessment 2013: Arctic Ocean acidification*, Oslo, Arctic Monitoring and
495 Assessment Programme.
- 496 Anderson LG, Björk G, Jutterström S, Pipko I, Shakhova N, Semiletov I, Wåhlström I (2011) East
497 Siberian Sea, an Arctic region of very high biogeochemical activity. *Biogeosciences*, **8**, 1745-
498 1754.
- 499 Anderson MJ (2001) A new method for non-parametric multivariate analysis of variance. *Austral*
500 *Ecology*, **26**, 32-46.
- 501 Arnkværn G, Daase M, Eiane K (2005) Dynamics of coexisting *Calanus finmarchicus*, *Calanus glacialis*
502 and *Calanus hyperboreus* populations in a high-Arctic fjord. *Polar Biology*, **28**, 528-538.
- 503 Azetsu-Scott K, Clarke A, Falkner K *et al.* (2010) Calcium carbonate saturation states in the waters of
504 the Canadian Arctic Archipelago and the Labrador Sea. *Journal of Geophysical Research:*
505 *Oceans*, **115**, C11021.
- 506 Beaugrand G, Brander KM, Lindley JA, Souissi S, Reid PC (2003) Plankton effect on cod recruitment in
507 the North Sea. *Nature*, **426**, 661-664.
- 508 Bellerby RGJ, Silyakova A, Nondal G, Slagstad D, Czerny J, De Lange T, Ludwig A (2012) Marine
509 carbonate system evolution during the EPOCA Arctic pelagic ecosystem experiment in the
510 context of simulated Arctic ocean acidification. *Biogeosciences Discuss.*, **2012**, 15541-15565.
- 511 Bopp L, Resplandy L, Orr JC *et al.* (2013) Multiple stressors of ocean ecosystems in the 21st century:
512 projections with CMIP5 models. *Biogeosciences*, **10**, 6225-6245.
- 513 Brown JH, Gillooly JF, Allen AP, Savage VM, West GB (2004) Toward a metabolic theory of ecology.
514 *Ecology*, **85**, 1771-1789.
- 515 Båmstedt U, Gifford DJ, Irigoien X, Atkinson DE, Roman MR (2000) Feeding. In: *Zooplankton*
516 *methodology handbook*. (eds Harris R, Wiebe PH, Lenz J, Skjoldal HR, Huntley ME) pp Page.
517 Oxford, Academic Press.

518 Caldeira K, Wickett ME (2005) Ocean model predictions of chemistry changes from carbon dioxide
519 emissions to the atmosphere and ocean. *Journal of Geophysical Research-Oceans*, **110**,
520 C09S04.

521 Calosi P, De Wit P, Thor P, Dupont S (2016) Will Life find a Way? Evolution of Marine Species Under
522 Global Change. *Evolutionary Applications*, n/a-n/a.

523 Calosi P, Melatunan S, Turner LM *et al.* (2017) Regional adaptation defines sensitivity to future ocean
524 acidification. *Nature Communications*, **8**, 13994.

525 Cao L, Caldeira K, Jain AK (2007) Effects of carbon dioxide and climate change on ocean acidification
526 and carbonate mineral saturation. *Geophysical Research Letters*, **34**, L05607.

527 Carroll SP, Hendry AP, Reznick DN, Fox CW (2007) Evolution on ecological time-scales. *Functional*
528 *Ecology*, **21**, 387-393.

529 Castonguay M, Plourde S, Robert D, Runge JA, Fortier L (2008) Copepod production drives
530 recruitment in a marine fish. *Canadian Journal of Fisheries and Aquatic Sciences*, **65**, 1528-
531 1531.

532 Clark RA, Frid CLJ, Nicholas KR (2003) Long-term, predation-based control of a central-west North Sea
533 zooplankton community. *ICES Journal of Marine Science*, **60**, 187-197.

534 Clarke A (1980) A reappraisal of the concept of metabolic cold adaptation in polar marine
535 invertebrates. *Biological Journal of the Linnean Society*, **14**, 77-92.

536 David JR, Gibert P, Gravot E, Petavy G, Morin J-P, Karan D, Moreteau B (1997) Phenotypic plasticity
537 and developmental temperature in *Drosophila*: Analysis and significance of reaction norms of
538 morphometrical traits. *Journal of Thermal Biology*, **22**, 441-451.

539 De Wit P, Dupont S, Thor P (2015) Selection on oxidative phosphorylation and ribosomal structure as
540 a multigenerational response to ocean acidification in the common copepod *Pseudocalanus*
541 *acuspes*. *Evolutionary Applications*, 1112-1123.

542 Dickson AG, Millero FJ (1987) A comparison of the equilibrium constants for the dissociation of
543 carbonic acid in seawater media. *Deep Sea Research Part A. Oceanographic Research Papers*,
544 **34**, 1733-1743.

545 Dupont S, Dorey N, Thorndyke M (2010) What meta-analysis can tell us about vulnerability of marine
546 biodiversity to ocean acidification? *Estuarine, Coastal and Shelf Science*, **89**, 182-185.

547 Dupont S, Pörtner HO (2013) Get ready for ocean acidification. *Nature*, **498**, 429-429.

548 Fabry VJ, McClintock JB, Mathis JT, Grebmeier JM (2009) Ocean acidification at high latitudes: The
549 bellweather. *Oceanography*, **22**, 160-171.

550 Fabry VJ, Seibel BA, Feely RA, Orr JC (2008) Impacts of ocean acidification on marine fauna and
551 ecosystem processes. *ICES Journal of Marine Science*, **65**, 414-432.

552 Fransson A, Chierici M, Anderson LG, Bussmann I, Kattner G, Peter Jones E, Swift JH (2001) The
553 importance of shelf processes for the modification of chemical constituents in the waters of
554 the Eurasian Arctic Ocean: implication for carbon fluxes. *Continental Shelf Research*, **21**, 225-
555 242.

556 Fransson A, Chierici M, Hop H, Findlay HS, Kristiansen S, Wold A (2016) Late winter-to-summer
557 change in ocean acidification state in Kongsfjorden, with implications for calcifying
558 organisms. *Polar Biology*, 1-17.

559 Freese D, Niehoff B, Søreide JE, Sartoris FJ (2015) Seasonal patterns in extracellular ion
560 concentrations and pH of the Arctic copepod *Calanus glacialis*. *Limnology and Oceanography*,
561 **60**, 2121-2129.

562 Frost BW (1972) Effect of size and concentration of food particles on the feeding behaviour of the
563 marine planktonic copepod *Calanus finmarchicus*. *Limnology and Oceanography*, **17**, 805-
564 815.

565 Grainger EH (1965) Zooplankton from the Arctic Ocean and adjacent Canadian waters. *Journal of the*
566 *Fisheries Research Board of Canada*, **22**, 543-564.

567 Halvorsen E (2015) Significance of lipid storage levels for reproductive output in the Arctic copepod
568 *Calanus hyperboreus*. *Marine Ecology Progress Series*, **540**, 259-265.

569 Hansen MO, Nielsen TG, Stedmon CA, Munk P (2012) Oceanographic regime shift during 1997 in
570 Disko Bay, Western Greenland. *Limnology and Oceanography*, **57**, 634-644.

571 Havenhand J, Dupont S, Quinn GP (2010) Designing ocean acidification experiments to maximise
572 inference. In: *Guide for best practices for ocean acidification research and data reporting*.
573 (eds Riebesell U, Fabry VJ, Hansson L, Gattuso JP) pp Page. Brussels, European Commission.

574 Hellberg ME (2009) Gene flow and isolation among populations of marine animals. *Annual Review*
575 *of Ecology Evolution and Systematics*, **40**, 291-310.

576 Hildebrandt N, Niehoff B, Sartoris FJ (2014) Long-term effects of elevated CO₂ and temperature on
577 the Arctic calanoid copepods *Calanus glacialis* and *C. hyperboreus*. *Marine Pollution Bulletin*,
578 **80**, 59-70.

579 Hildebrandt N, Sartoris FJ, Schulz KG, Riebesell U, Niehoff B (2016) Ocean acidification does not alter
580 grazing in the calanoid copepods *Calanus finmarchicus* and *Calanus glacialis*. *ICES Journal of*
581 *Marine Science*, **73**, 927-936.

582 Hoegh-Guldberg O, Cai R, Poloczanska ES *et al.* (2014) The Ocean. In: *Climate Change 2014: Impacts,*
583 *Adaptation, and Vulnerability. Part B: Regional Aspects. Contribution of Working Group II to*
584 *the Fifth Assessment Report of the Intergovernmental Panel of Climate Change*. (eds Barros
585 VR, Field CB, Dokken DJ, Mastrandrea MD, Mach KJ, Bilir TE, Chatterjee M, Ebi KL, Estrada YO,
586 Genova RC, Girma B, Kissel ES, Levy AN, MacCracken S, Mastrandrea PR, White LL) pp Page.
587 Cambridge, United Kingdom and New York, NY, USA, Cambridge University Press.

588 Ikeda T, Kanno Y, Ozaki K, Shinada A (2001) Metabolic rates of epipelagic marine copepods as a
589 function of body mass and temperature. *Marine Biology*, **139**, 587-596.

590 IPCC (2013) *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the*
591 *Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, Cambridge, UK
592 and New York, USA, Cambridge University Press.

593 Isari S, Zervoudaki S, Saiz E, Pelejero C, Peters J (2015) Copepod vital rates under CO₂-induced
594 acidification: a calanoid species and a cyclopoid species under short-term exposures. *Journal*
595 *of Plankton Research*.

596 Jutterström S, Anderson LG (2010) Uptake of CO₂ by the Arctic Ocean in a changing climate. *Marine*
597 *Chemistry*, **122**, 96-104.

598 Klekowski RZ, Weslawski JM (1991) *Atlas of the marine fauna of Southern Spitsbergen*.

599 Last JM (1980) *The food of twenty species of fish larvae in the west-central North Sea*, Lowestoft (UK),
600 Ministry of Agriculture, Fisheries and Food.

601 Li W, Gao K (2012) A marine secondary producer respire and feeds more in a high CO₂ ocean.
602 *Marine Pollution Bulletin*, **64**, 699-703.

603 Mackas DL, Goldblatt R, Lewis AG (1998) Interdecadal variation in developmental timing of
604 *Neocalanus plumchrus* populations at Ocean Station P in the subarctic North Pacific.
605 *Canadian Journal of Fisheries and Aquatic Sciences*, **55**, 1878-1893.

606 Madsen SD, Nielsen TG, Hansen BW (2001) Annual population development and production by
607 *Calanus finmarchicus*, *C. glacialis* and *C. hyperboreus* in Disko Bay, western Greenland.
608 *Marine Biology*, **139**, 75-93.

609 Mauchline J (1998) *The biology of calanoid copepods*, San Diego, Academic Press.

610 Mehrbach C, Culbertson CH, Hawley JE, Pytkowicz RM (1973) Measurement of the apparent
611 dissociation constant of carbonic acid in seawater at atmospheric pressure. *Limnology and*
612 *Oceanography*, **18**, 897-907.

613 Møller EF, Nielsen TG, Richardson K (2006) The zooplankton community in the Greenland Sea:
614 Composition and role in carbon turnover. *Deep-Sea Research Part I-Oceanographic Research*
615 *Papers*, **53**, 76-93.

616 Möllmann C, Kornilovs G, Fetter M, Koster FW, Hinrichsen HH (2003) The marine copepod,
617 *Pseudocalanus elongatus*, as a mediator between climate variability and fisheries in the
618 Central Baltic Sea. *Fisheries Oceanography*, **12**, 360-368.

619 Nielsen TG, Kjellerup S, Smolina I, Hoarau G, Lindeque P (2014) Live discrimination of *Calanus*
620 *glacialis* and *C. finmarchicus* females: can we trust phenological differences? Marine Biology,
621 **161**, 1299-1306.

622 Omori M, Ikeda T (1984) *Methods in marine zooplankton ecology*, New York, Wiley.

623 Pedersen SA, Hakedal OJ, Salaberria I *et al.* (2014) Multigenerational exposure to ocean acidification
624 during food limitation reveals consequences for copepod scope for growth and vital rates.
625 Environmental Science & Technology, **48**, 12275-12284.

626 Peijnenburg KTCA, Goetze E (2013) High evolutionary potential of marine zooplankton. Ecology and
627 Evolution, **3**, 2765-2781.

628 Pespeni MH, Sanford E, Gaylord B *et al.* (2013) Evolutionary change during experimental ocean
629 acidification. Proc.Natl.Acad.Sci., **110**, 6937-6942.

630 Peterson BJ, Holmes RM, McClelland JW *et al.* (2002) Increasing River Discharge to the Arctic Ocean.
631 Science, **298**, 2171-2173.

632 Pierrot D, Lewis E, Wallace DWR (2006) *MS Excel program developed for CO2 system calculations.*
633 *ORNL/CDIAC-105a.*, Oak Ridge, Tennessee, Carbon Dioxide Information Analysis Center, Oak
634 Ridge National Laboratory, US Department of Energy.

635 Rastrick SP, Whiteley NM (2011) Congeneric amphipods show differing abilities to maintain
636 metabolic rates with latitude. Physiological and Biochemical Zoology, **84**, 154-165.

637 Raymond PA, McClelland JW, Holmes RM *et al.* (2007) Flux and age of dissolved organic carbon
638 exported to the Arctic Ocean: A carbon isotopic study of the five largest arctic rivers. Global
639 Biogeochemical Cycles, **21**, GB4011.

640 Rey-Rassat C, Irigoien X, Harris R, Head R, Carlotti F (2002) Growth and development of *Calanus*
641 *helgolandicus* reared in the laboratory. Marine Ecology Progress Series, **238**, 125-138.

642 Riebesell U, Fabry VJ, Hansson L, Gattuso JP (2010) *Guide to best practice for research for ocean*
643 *acidification and data reporting*, Luxembourg, Publications Office of the European Union.

644 Riisgaard K, Nielsen TG, Hansen PJ (2015) Impact of elevated pH on succession in the Arctic spring
645 bloom. Marine Ecology Progress Series, **530**, 63-75.

646 Rocha FB, Klaczko LB (2012) Connecting the dots of nonlinear reaction norms unravels the threads of
647 genotype-environment interaction in *Drosophila*. Evolution, **66**, 3404-3416.

648 Runge JA, Castonguay M, De Lafontaine Y, Ringuette M, Beaulieu JL (1999) Covariation in climate,
649 zooplankton biomass and mackerel recruitment in the southern Gulf of St Lawrence.
650 Fisheries Oceanography, **8**, 139-149.

651 Runge JA, Fields DM, Thompson CRS *et al.* (2016) End of the century CO2 concentrations do not have
652 a negative effect on vital rates of *Calanus finmarchicus*, an ecologically critical planktonic
653 species in North Atlantic ecosystems. ICES Journal of Marine Science, **73**, 937-950.

654 Sanford E, Kelly MW (2010) Local Adaptation in Marine Invertebrates. Annual Review of Marine
655 Science, **3**, 509-535.

656 Sibly RM, Calow P (1986) *Physiological ecology of animals - an evolutionary approach*, Oxford,
657 Blackwell Scientific publications.

658 Small DP, Calosi P, Boothroyd D, Widdicombe S, Spicer JI (2015) Stage-specific changes in
659 physiological and life-history responses to elevated temperature and pCO2 during the larval
660 development of the European Lobster *Homarus gammarus* (L.). Physiological and
661 Biochemical Zoology, **88**, 494-507.

662 Steinacher M, Joos F, Frölicher TL, Plattner GK, Doney SC (2009) Imminent ocean acidification in the
663 Arctic projected with the NCAR global coupled carbon cycle-climate model. Biogeosciences,
664 **6**, 515-533.

665 Strickland JD, Parsons TR (1972) A practical handbook of seawater analysis. Journal of the Fisheries
666 Research Board of Canada, **167**, 310.

667 Stumpp M, Hu M, Casties I, Saborowski R, Bleich M, Melzner F, Dupont S (2013) Digestion in sea
668 urchin larvae impaired under ocean acidification. Nature Clim.Change, **3**, 1044-1049.

669 Sunday JM, Calosi P, Dupont S, Munday PL, Stillman JH, Reusch TBH (2014) Evolution in an acidifying
670 ocean. Trends in Ecology & Evolution, **29**, 117-125.

- 671 Thoisen C, Riisgaard K, Lundholm N, Nielsen TG, Hansen PJ (2015) Effect of acidification on an Arctic
672 phytoplankton community from Disko Bay, West Greenland. *Marine Ecology Progress Series*,
673 **520**, 21-34.
- 674 Thor P, Bailey A, Halsband C, Guscelli E, Gorokhova E, Fransson A (2016) Seawater pH predicted for
675 the year 2100 affects the metabolic response to feeding in copepodites of the Arctic copepod
676 *Calanus glacialis*. *PLoS ONE*, **11**, e0168735.
- 677 Thor P, Dupont S (2015) Transgenerational effects alleviate severe fecundity loss during ocean
678 acidification in a ubiquitous planktonic copepod. *Global Change Biology*, **21**, 2261-2271.
- 679 Thor P, Koski M, Tang KW, Jónasdóttir SH (2007) Supplemental effects of diet mixing on absorption of
680 ingested organic carbon in the marine copepod *Acartia tonsa*. *Marine Ecology Progress
681 Series*, **331**, 131-138.
- 682 Thor P, Nielsen TG, Tiselius P (2008) Mortality rates of epipelagic copepods in the post-spring bloom
683 period in the Disko Bay, western Greenland. *Marine Ecology Progress Series*, **359**, 151-160.
- 684 Thor P, Nielsen TG, Tiselius P *et al.* (2005) Post spring bloom community structure of pelagic
685 copepods in the Disko Bay, Western Greenland. *Journal of Plankton Research*, **27**, 341-356.
- 686 Thor P, Oliva EO (2015) Ocean acidification elicits different energetic responses in an Arctic and a
687 boreal population of the copepod *Pseudocalanus acuspes*. *Marine Biology*, **162**, 799-807.
- 688 Thor P, Wendt I (2010) Functional response of carbon absorption efficiency in the copepod *Acartia
689 tonsa* Dana. *Limnology and Oceanography*, **55**, 1779-1789.
- 690 Vargas CA, Lagos NA, Lardies MA *et al.* (2017) Species-specific responses to ocean acidification should
691 account for local adaptation and adaptive plasticity. *Nature Ecology & Evolution*, **1**, 0084.
- 692 Walther K, Anger K, Pörtner HO (2010) Effects of ocean acidification and warming on the larval
693 development of the spider crab *Hyas araneus* from different latitudes (54° vs. 79°N). *Marine
694 Ecology Progress Series*, **417**, 159-170.
- 695 Wassmann P, Kosobokova KN, Slagstad D *et al.* (2015) The contiguous domains of Arctic Ocean
696 advection: Trails of life and death. *Progress in Oceanography*, **139**, 42-65.
- 697 Whiteley NM (2011) Physiological and ecological responses of crustaceans to ocean acidification.
698 *Marine Ecology Progress Series*, **430**, 257-271.
- 699 Wittmann AC, Pörtner HO (2013) Sensitivities of extant animal taxa to ocean acidification. *Nature
700 Clim.Change*, **3**, 995-1001.
- 701 Wood HL, Sundell K, Almroth BC, Sköld HN, Eriksson SP (2016) Population-dependent effects of
702 ocean acidification. *Proceedings of the Royal Society of London B: Biological Sciences*, **283**.
- 703 Yamamoto-Kawai M, Mclaughlin FA, Carmack EC, Nishino S, Shimada K (2009) Aragonite
704 undersaturation in the Arctic Ocean: Effects of Ocean Acidification and Sea Ice Melt. *Science*,
705 **326**, 1098-1100.
- 706 Zervoudaki S, Frangoulis C, Giannoudi E, Krasakopoulou E (2014) Effects of low pH and raised
707 temperature on egg production, hatching and metabolic rates of a Mediterranean copepod
708 species (*Acartia clausi*) under oligotrophic conditions. *Mediterranean Marine Science*, **15**, 74-
709 83.

710

711

712 **Table 1**

713 Number of replicates *per* treatment combination: copepodite developmental stage (CIII, CIV,
 714 CV) of *Calanus glacialis* by nominal pH level according to our experimental design. When
 715 different, numbers preceding that slash refer to ingestion rate measurements and number
 716 following the slash refer to metabolic rate measurements. When only one value is indicated
 717 the number of replicates were equal. A total number of 153 samples were included in analyses
 718 of ingestion rate and a total of 170 in analyses of metabolic rates. By necessity the number of
 719 replicates varied with the number of copepodites available.

720

Copepodite stage	Location	Nominal pH								
		8.2	8.0	7.8	7.6	7.4	7.2	7.0	6.6	6.4
CIII	Kongsfjord	3/2	3	1/2	3	2/3	2	3	1	1
CIV	Kongsfjord	4	4	2	4/3	4	2/0	4/3	2	2
	Billefjord	4/3	4/3	2/4	3	4	3	4		
	Disko Bay	0/4	3/1	1	5	5	1	4	2/3	3/4
CV	Kongsfjord	4/3	4	2/1	3/4	4	2	4	2	1/2
	Billefjord	0/1	1	1	1	1	1	1		
	Disko Bay	0/3	6/8	2/5	4	5	1/5	3/5	2/3	2/3

721

722 **Table 2**

723 Mean \pm standard deviations of carbonate chemistry parameters during incubations. pH_{nom} is
 724 nominal pH treatment, pH_{T} is total hydrogen scale pH, A_{T} is total alkalinity, and $p\text{CO}_2$ is CO_2
 725 partial pressure. A_{T} was measured only once in the pH_{nom} 7.5 treatment once in the Billefjord
 726 population experiment.

727

pH_{nom}	T °C	S	pH_{T}	A_{T} $\mu\text{mol kg}^{-1}$	$p\text{CO}_2$ μatm
<i>Kongsfjord</i>					
8.1	6.1 ± 0.9	34.0 ± 0.1	8.012 ± 0.064	$2\,347 \pm 11$	450 ± 95
7.9	6.0 ± 0.7	34.0 ± 0.1	7.851 ± 0.062	$2\,351 \pm 14$	712 ± 134
7.7	5.9 ± 0.6	33.9 ± 0.1	7.618 ± 0.092	$2\,354 \pm 9$	$1\,213 \pm 346$
7.5	6.3 ± 0.8	34.0 ± 0.1	7.442 ± 0.088	$2\,353 \pm 21$	$1\,973 \pm 460$
7.3	6.4 ± 0.7	34.0 ± 0.1	7.318 ± 0.067	$2\,348 \pm 11$	$2\,414 \pm 308$
7.1	6.3 ± 0.8	34.0 ± 0.1	7.160 ± 0.063	$2\,353 \pm 5$	$3\,546 \pm 543$
6.9	6.4 ± 0.7	34.0 ± 0.1	6.998 ± 0.044	$2\,350 \pm 16$	$5\,132 \pm 526$
6.6	6.5 ± 0.5	34.0 ± 0.1	6.636 ± 0.050	$2\,337 \pm 8$	$11\,534 \pm 1\,368$
6.4	6.3 ± 0.4	34.0 ± 0.1	6.445 ± 0.039	$2\,332 \pm 6$	$18\,567 \pm 2\,163$
<i>Billefjord</i>					
8.1	6.2 ± 0.7	34.0 ± 0.1	8.041 ± 0.056		446 ± 93
7.9	6.5 ± 0.7	34.0 ± 0.1	7.851 ± 0.033		683 ± 49
7.7	6.5 ± 0.4	34.0 ± 0.1	7.644 ± 0.047		$1\,119 \pm 178$
7.5	6.9 ± 0.7	34.0 ± 0.1	7.497 ± 0.034	$2\,322 \pm 3$	$1\,536 \pm 154$
7.3	6.5 ± 0.4	34.0 ± 0.1	7.337 ± 0.036		$2\,319 \pm 257$
7.1	6.5 ± 0.5	34.0 ± 0.1	7.180 ± 0.043		$3\,336 \pm 392$
6.9	6.4 ± 0.4	34.1 ± 0.1	7.036 ± 0.041		$4\,526 \pm 499$
<i>Disko Bay</i>					
8.1	3.9 ± 0.3	34.4 ± 0.1	8.001 ± 0.059	$2\,280 \pm 0$	436 ± 64
7.9	3.9 ± 0.4	34.4 ± 0.1	7.805 ± 0.050	$2\,286 \pm 7$	721 ± 91
7.7	3.6 ± 0.6	34.4 ± 0.1	7.627 ± 0.046	$2\,293 \pm 0$	$1\,112 \pm 128$
7.5	3.9 ± 0.7	34.4 ± 0.1	7.431 ± 0.056	$2\,287 \pm 7$	$1\,774 \pm 250$
7.3	3.6 ± 0.5	34.6	7.262 ± 0.036	$2\,287 \pm 7$	$2\,642 \pm 233$
7.1	3.5 ± 0.5	34.6	7.099 ± 0.020	$2\,293 \pm 0$	$3\,865 \pm 194$
6.9	3.6 ± 0.6	34.6	6.920 ± 0.040	$2\,287 \pm 7$	$5\,878 \pm 517$
6.6	4.5 ± 0.8	34.6	6.562 ± 0.037	$2\,280 \pm 0$	$13\,325 \pm 1\,130$
6.4	4.1 ± 0.8	34.6	6.403 ± 0.077	$2\,280 \pm 0$	$19\,456 \pm 3\,521$

728

729 **Table 3**

730 Ingestion rate reaction norms of *Calanus glacialis* copepodite stage CIV. Results of the first
 731 order regression model, $IR = \bar{IR} + g(pH_T - \bar{pH}_T)$ (David *et al.*, 1997), where \bar{IR} is mean
 732 ingestion rate, g is the slope, and \bar{pH}_T is mean pH_T .

Stage	Location	\bar{IR} $\mu\text{gC } \mu\text{gC}^{-1} \text{ d}^{-1}$	\bar{pH}_T	g $\times 10^{-3}$	r^2	P
CIII	Kongsfjord	0.1006	7.30	17.6	0.05	0.369 ⁺
CIV	Kongsfjord	0.0474	7.30	39.4	0.41	<0.001
	Billefjord	0.0398	7.62	28.0	0.19	0.031
	Disko Bay	0.0456	7.30	7.13	0.04	0.323
CV	Kongsfjord	0.0271	7.24	13.8	0.11	0.111
	Billefjord	0.0234	7.51	-2.23	0.02	0.808
	Disko Bay	0.0121	7.31	-0.91	0.02	0.540

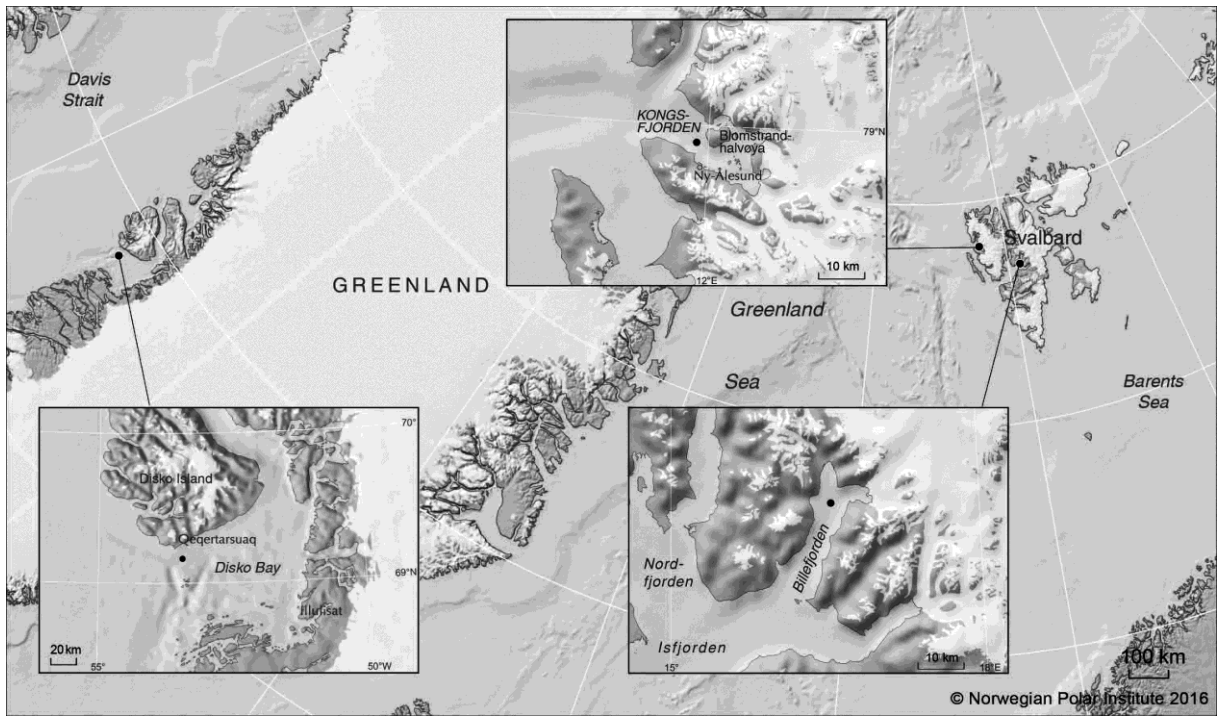
733 ⁺ Ingestion rates of CIIIs were best fitted with the second order regression
 734 model (see text).

735 **Table 4**

736 Metabolic rate reaction norms of *Calanus glacialis* copepodite stage CIV. Results of the first
 737 order regression model, $\dot{M}O_2 = \bar{M}O_2 + g(pH_T - \bar{p}H_T)$ (David *et al.*, 1997), where $\bar{M}O_2$ is
 738 mean metabolic rate, g is the slope, and $\bar{p}H_T$ is mean pH_T.

Stage	Location	$\bar{M}O_2$ μgC μgC ⁻¹ d ⁻¹	$\bar{p}H_T$	g x10 ⁻³	r ²	P
CIII	Kongsfjord	0.0210	7.29	15.1	0.16	0.080
CIV	Kongsfjord	0.0206	7.30	-6.81	0.15	0.043
	Billefjord	0.0254	7.62	-10.2	0.23	0.014
	Disko Bay	0.0258	7.31	-2.92	0.03	0.359
CV	Kongsfjord	0.0170	7.30	2.51	0.04	0.354
	Billefjord	0.0149	7.62	-6.18	0.12	0.456
	Disko Bay	0.0231	7.31	-3.81	0.04	0.236

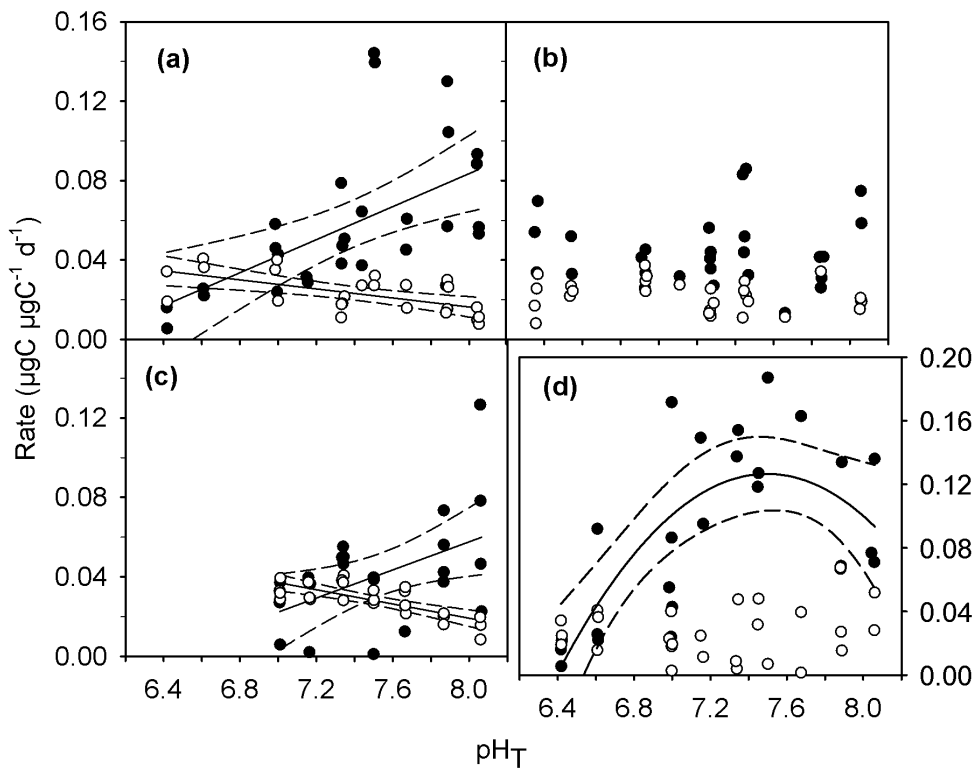
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741 Figure 1. Study sites in Kongsfjord, Billefjord (Svalbard), and Disko Bay (West Greenland).

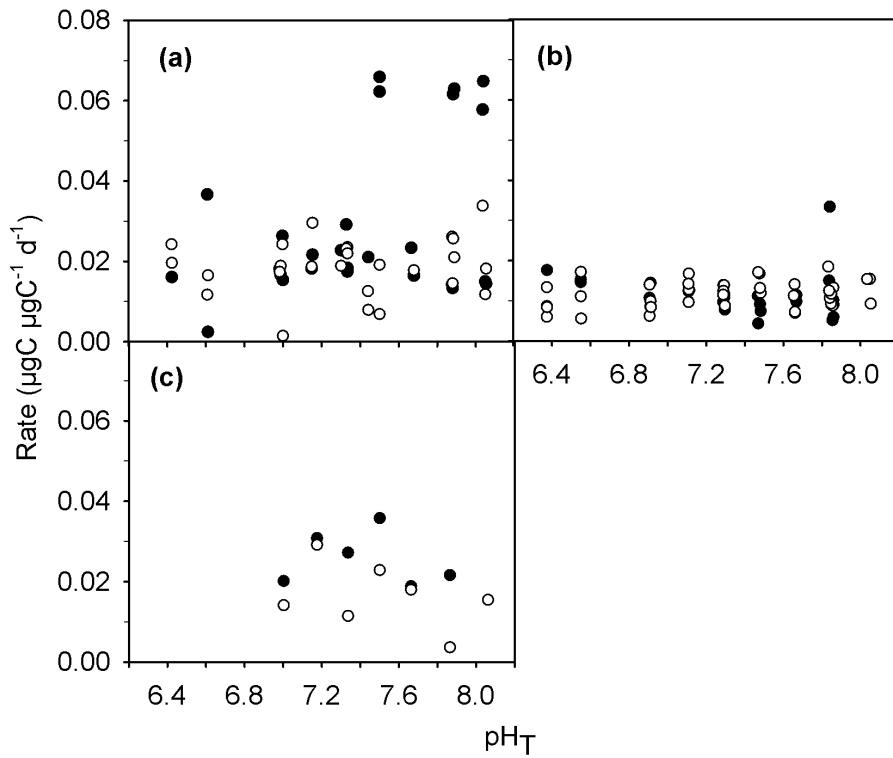
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744 Figure 2. *Calanus glacialis* copepodite stages CIII and CIV. Ingestion rates (filled circles) and
 745 metabolic rates (open circles) vs. seawater pH_T in the three populations. a) Kongsfjord CIVs,
 746 b) Disko Bay CIVs, c) Billefjord CVs, and d) Kongsfjord CIIIs. Lines depict first or second
 747 order reaction norms. Solid lines show predicted values and hatched lines show 95%
 748 confidence limits. Reaction norm parameters and statistics are shown in Tables 3 and 4.

749



750

751 Figure 3. *Calanus glacialis* copepodite stage CV. Ingestion rates (filled circles) and metabolic
 752 rates (open circles) vs. seawater pH_T in the three populations investigated. a) Kongsfjord, b)
 753 Disko Bay, and c) Billefjord.