

Solar Flare Effects on 150-km Echoes Observed Over Jicamarca: WACCM-X Simulations

N. M. Pedatella^{1,2}, J. L. Chau³, J. Vierinen⁴, L. Qian¹, P. Reyes⁵, E. Kudeki⁶, G. Lehmacher⁷, and M. Oppenheim⁸

¹High Altitude Observatory, National Center for Atmospheric Research, Boulder, CO, USA.

²COSMIC Program Office, University Center for Atmospheric Research, Boulder, CO, USA.

³Leibniz Institute of Atmospheric Physics, Rostock University, Kühlungsborn, Germany.

⁴Department of Physics and Technology, University of Tromsø, Tromsø, Norway

⁵Center for Geospace Studies, SRI International, Menlo Park, California, USA

⁶Department of Electrical and Computer Engineering, University of Illinois at Urbana-Champaign,

⁷Department of Physics and Astronomy, University of Clemson, Clemson, SC, USA.

⁸Center for Space Physics, Boston University, Boston, MA, USA.

Key Points:

- There is a good agreement between observed morphology of 150-km echoes and simulated electron densities during a solar flare.
- The results support the hypothesis that layering of 150-km echoes is connected to electron densities.
- Decrease in vertical plasma drift during the solar flare can be attributed to changes in E-region conductivity.

Abstract

Jicamarca Radio Observatory observations and Whole Atmosphere Community Climate Model with thermosphere-ionosphere eXtension (WACCM-X) simulations are used to investigate the effects of the September 7, 2005 X-17 solar flare on 150-km echoes, electron densities, and vertical plasma drifts. The solar flare produces a remarkably similar response in the observed 150-km echoes and simulated electron densities. The results provide additional evidence of the relationship between the background electron density and the layering structure that is seen in 150-km echoes. The simulations also capture a similar rapid decrease in vertical plasma drift velocity that is seen in the observations. The simulated change in vertical plasma drift is, however, weaker than the observed decrease at the longitude of Jicamarca, though it is stronger east of Jicamarca. The effect of the solar flare on the vertical plasma drifts is **primarily** attributed to changes in conductivity due to the enhanced ionization during the solar flare.

1 Introduction

Despite being first observed over 50-years ago (Balsley, 1964), the source of the enhanced VHF radar echoes that are observed near 150 km remains unexplained. These so-called 150-km echoes have subsequently been observed in the equatorial ionosphere at multiple longitudinal locations (J. L. Chau & Kudeki, 2006; Kudeki et al., 1998; de Paula & Hysell, 2004; Tsunoda & Ecklund, 2008; Choudhary et al., 2004; A. K. Patra et al., 2008). The 150-km echoes are observed nearly every day, and they are thus a ubiquitous feature of the equatorial ionosphere. The characteristics of the 150-km echoes have been well documented by observations. These characteristics include their occurrence only during the daytime, a necklace-like shape with descending structures prior to noon and ascending structures after noon, as well as the formation of distinct layers. Moreover, the majority of 150-km echoes are a manifestation of naturally enhanced incoherent scatter echoes (e.g., J. L. Chau, 2004; J. Chau et al., 2009; J. L. Chau & Kudeki, 2013). Some of these features have led to the hypothesis that the 150-km echoes are due to photoelectrons (e.g., Oppenheim & Dimant, 2016). The connection to photoelectrons is further supported by the absence of 150-km echoes during the January 2010 solar eclipse (A. K. Patra et al., 2011), as well as their modification by solar flares (Reyes, 2012).

The photoelectron origin of the 150-km echoes cannot fully explain the formation of several distinct layers. There does, however, appear to be a connection between the electron density, the temporal and altitudinal structure of the layers, and the gaps that form between layers (e.g., J. L. Chau et al., 2009; Reyes, 2017). For example, Reyes (2017) show a close correspondence between short-period (~ 5 -10 minute) fluctuations in electron densities and the gaps in the 150-km echoes. This has led to the suggestion that the gaps between layers may form at distinct plasma frequencies (e.g., G. Lehmacher et al., 2018). However, A. Patra et al. (2017) recently disputed the connection between photoelectrons and 150-km echoes. They found an inverse relationship between 150-km echo power and EUV flux, and hypothesized that neutral dynamics play an important role in the formation of 150-km echoes. Though neutral dynamics may contribute to the 150-km echoes, they would not fully explain features such as the daytime only occurrence, narrow spectral widths, solar eclipse, and solar flare effects. These features suggest that photoelectrons contribute to the formation of 150-km echoes.

Nonetheless, the connection between 150-km echoes, photoelectrons, and electron densities has yet to be fully explored. Understanding the relationship between 150-km echoes and electron densities would provide an additional step towards developing a complete theory to explain the 150-km echoes, improving our understanding of the equatorial ionosphere. If a relationship between electron density and the power striations can be determined, 150-km echoes could also provide a high signal-to-noise ratio radar tar-

70 get for accurately measuring the electron density between the E-region and the F-region
71 of the ionosphere.

72 The main objective of the present study is to further investigate the connection be-
73 tween electron densities and the 150-km echo layers. This is done through a compari-
74 son of Jicamarca Radio Observatory (JRO) observations of 150-km echoes during the
75 September 7, 2005 solar flare with electron densities simulated in the Whole Atmosphere
76 Community Climate Model with thermosphere-ionosphere eXtension (WACCM-X). The
77 simulated electron densities closely follow the observed layering structure of the 150-km
78 echoes, supporting the close connection between the electron densities and the gaps that
79 form between the 150-km echo layers. We further investigate the change in vertical plasma
80 drift velocity during the solar flare, and find that this is likely related to a rapid change
81 in the conductivity that occurs during the solar flare.

82 2 JRO Observations

83 The JRO observations were taken as part of a MST-ISR experiment, which is an
84 experiment designed to observe the Mesosphere, Stratosphere and Troposphere (MST)
85 at the same time as the ionosphere in quasi thermal equilibrium via the Incoherent Scatter
86 Radar (ISR) mode (e.g., G. A. Lehmacher et al., 2009, 2019). The MST mode allows
87 the observation from 0 km to 200 km, while the ISR mode measures from 200 km to 900
88 km in altitude. Although the mesosphere stops at around 100 km, the MST mode has
89 proven to also be useful in the study of coherent scattering from ionospheric irregular-
90 ities, such as those coming from 150-km echoes (e.g., Kudeki & Fawcett, 1993; J. L. Chau
91 & Kudeki, 2006). The MST-ISR mode is realized by interleaving sequences of pulses with
92 different repetition, pulse width and pulse coding. In the case of the MST part, 20 con-
93 secutive pulses with 1.33 ms (or ~ 200 km) interpulse period (IPP) and 64 baud com-
94plementary codes pulses with a total width of 64 μ s (or 9.6 km) are transmitted. In the
95 case of the ISR part, 2 Barker-3 coded pulses with a total width of 300 μ s (or 45 km)
96 and an IPP of 6.66 ms (or ~ 1000 km) were transmitted. **The nominal range reso-**
97 **lution of the observations is 150 m.**

98 These pulse sequences were transmitted simultaneously on four different beam po-
99 sitions (North, East, South and West), taking advantage of the modular and polariza-
100 tion features of JRO. Two transmitters of 1 MW peak power each, were combined be-
101 fore feeding all four beams simultaneously, i.e., on each beam 500 kW peak power was
102 transmitted. In this work we present the results of September 7, 2005 only from the MST
103 part of the West beam (-87.68° azimuth, 87.52° elevation), which is the beam pointing
104 the closest to perpendicular to the Earth's magnetic field \mathbf{B} (beam gain peak $\sim 0.8^\circ$ from
105 perpendicular to \mathbf{B} and elongated in the North-South direction with a beam width of
106 $\sim 1.4^\circ$) at 150 km at the time of the experiment. More details of the JRO modes, sig-
107 nal processing, other solar flare effects, and other events can be found in Reyes (2012).

108 3 WACCM-X

109 Model simulations are performed in WACCM-X version 2.0 (H.-L. Liu et al., 2018).
110 WACCM-X extends from the surface to the upper thermosphere (4.1×10^{-10} hPa, ~ 500 -
111 700 km depending on solar activity), and has a resolution of 1.9° in latitude, 2.5° in lon-
112 gitude, and 0.25 scale heights above the stratosphere. Up to the lower thermosphere, WACCM-
113 X is based on the Community Atmosphere Model (CAM) version 4 (Neale et al., 2013)
114 and Whole Atmosphere Community Climate Model (WACCM) version 4 (Marsh et al.,
115 2013). Upper atmospheric processes, including the transport of O^+ , self-consistent iono-
116 spheric electrodynamics, and energetics included in WACCM-X are primarily based on
117 the Thermosphere-Ionosphere-Electrodynamics General Circulation Model (TIE-GCM)
118 (Roble et al., 1988; Richmond et al., 1992). H.-L. Liu et al. (2018) and J. Liu et al. (2018)
119 provide a detailed description and validation, respectively, of WACCM-X version 2.0

For the model simulations in the present study, the specified dynamics approach (Smith et al., 2017) is used to constrain the lower atmosphere meteorology up to 50 km to the National Aeronautics and Space Administration (NASA) Modern Era Retrospective Analysis for Research and Applications version 2 (MERRA-2) (Gelaro et al., 2017). Geomagnetic forcing is incorporated by imposing the Heelis empirical convection pattern at high latitudes (Heelis et al., 1982), which is driven by the 3-hr geomagnetic K_p index. The Flare Irradiance Spectral Model (FISM) (Chamberlin et al., 2008) provides the solar spectral irradiance for the solar flare that occurred on September 7, 2005. FISM is an empirical model that uses observational data from Geostationary Operational Environmental Satellite (GOES) X-Ray Sensor (XRS), Thermosphere Ionosphere Mesosphere Energetics and Dynamics (TIMED) Solar Extreme Ultraviolet Experiment (SEE), and Solar Radiation and Climate Experiment (SORCE) Solar Stellar Irradiance Comparison Experiment (SOLSTICE) to estimate the solar irradiance at wavelengths from 0.1 to 190 nm at 60 s temporal resolution. FISM is thus able to capture the solar irradiance variability during solar flares at wavelengths that directly impact the ionosphere and thermosphere, which includes the soft X-rays (0.1-10 nm) and extreme ultraviolet (EUV, 10-121.6 nm). Previous studies have demonstrated that the solar flare irradiance information provided by FISM is suitable for studying the effects of solar flares in the mesosphere, thermosphere, and ionosphere (e.g., Qian et al., 2011; Pettit et al., 2018).

4 Results and Discussion

4.1 Flare Impact on 150-km Echoes and Electron Density

The signal to noise ratio (SNR) observed by JRO on September 7, 2005 is shown in Figure 1a. The X-ray flux observed by GOES XRS is shown in Figure 1c. An X-17 solar flare began at 17:17 UT, reached its maximum intensity at 17:40 UT, and the solar irradiance returned to nominal levels over the next \sim 1 hour. Prior to the solar flare, the characteristic behavior of 150-km echoes is observed, with gradually descending layers of enhanced SNR that are 5-10 km thick. The enhanced SNR layers are separated by gaps that are on the order of a kilometer thick. The layers descend rapidly in altitude beginning around 17:30 UT, which corresponds to the **time when the X-ray enhancement was observed by the GOES satellite**. After the flare, around 18:00 UT, the layers initially rise rapidly, though the rate of ascent slows over the following hour. The vertical thickness of the layers also appears to be changed by the solar flare, with the layers being narrower following the solar flare.

The corresponding electron densities simulated by WACCM-X are shown in Figure 1b. Note that the WACCM-X results have been shifted later by 5 minutes to be more consistent with the observations. This corresponds to the model time step, as well as the solar flare forcing input, so we consider a 5 minute offset to not be a significant discrepancy between the timing of the solar flare effects in the observations and simulations. We also note that the coarse (relative to solar flare time scales) time step of WACCM-X may tend to smooth the model response to the solar flare. Contours of constant electron density in the WACCM-X simulations exhibit many of the same features that are seen in the observations. Prior to the flare, the electron density contours can be seen to largely track the gaps and edges in the radar echoes. The exception being the smaller scale structures that are seen in the observations, which are attributed to gravity waves that are unresolved in WACCM-X. The consistency between the Jicamarca observations and WACCM-X simulations is especially apparent during the solar flare. In particular, both show a rapid descent in altitude beginning around 17:30 UT, followed by a more gradual ascent around 18:00 UT. The electron density contours are additionally more closely spaced following the flare, a feature consistent with the JRO SNR observations.

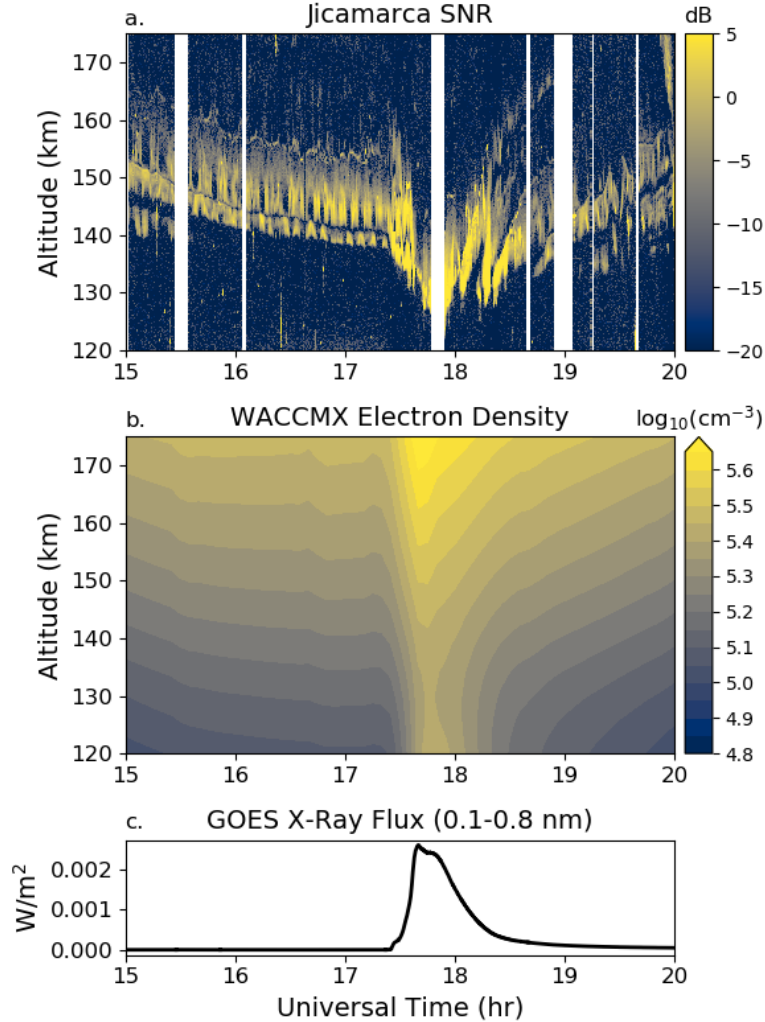


Figure 1. (a) JRO observed signal to noise ratio (SNR) during the September 7, 2005 solar flare. White areas indicate time periods without observations. (b) Electron densities simulated by WACCM-X at the location of Jicamarca, Peru. (c) Observed GOES x-ray flux for 0.1-0.8 nm.

169 To better illustrate the relationship between the 150-km echoes observed by JRO
 170 and the electron density simulated by WACCM-X, the two are plotted together in Fig-
 171 ure 2 for a shorter time interval around the solar flare. The remarkable agreement in the
 172 effect of the solar flare on contours of constant electron density and the structure of the
 173 150-km echo layers can be clearly seen in Figure 2. From Figure 2, it is apparent that
 174 the gaps in the 150-km echoes seem to follow electron density contours; however, the rea-
 175 son for this relationship is not yet known. With plasma-lines, there is a matching con-
 176 dition between plasma-frequency, radar wavelength, and suprathermal electron veloc-
 177 ity (photoelectrons and auroral secondary electrons), which results in electron density
 178 dependent plasma-line radar echo enhancements (Perkins et al., 1965). The fact that this
 179 also occurs for the 150-km echoes, points to a similar wave-particle interaction (Oppenheim
 180 & Dimant, 2016). An alternative possibility is that the observed layering is related to
 181 gyro-harmonics (G. Lehmacher et al., 2018), though this would not explain the forma-
 182 tion of multiple layers in the E-region because there are only two contours in the E-region
 183 where the electron density plasma frequency is an integer multiple of the gyro-frequency.

184 Thus, although the results demonstrate a close connection between electron density and
 185 the 150-km echo layers, the reason for this relationship remains unknown. Both of the
 186 previously mentioned hypotheses will be explored in detail in a future work, where com-
 187 parisons between JRO observations and WACCM-X simulations under nominal (i.e., non-
 188 flare) conditions will be considered.

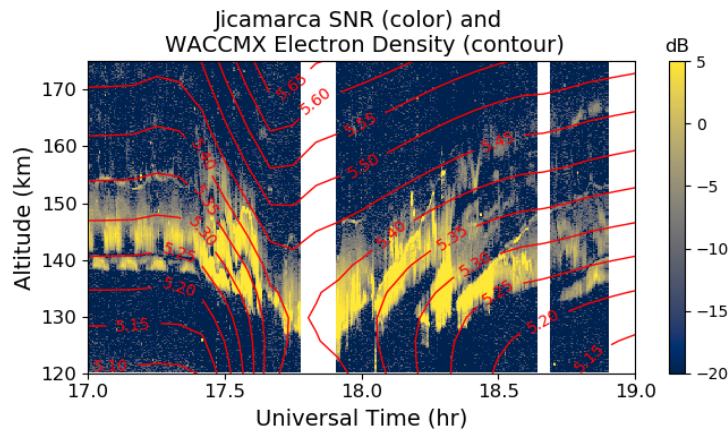


Figure 2. Observed signal to noise ratio (SNR) (colors), and WACCM-X electron densities in units of $\log_{10}\text{cm}^{-3}$ (contours) during the September 7, 2005 solar flare.

189 4.2 Flare Impact on Vertical Plasma Drifts

190 In addition to influencing the E-region electron densities and 150-km echoes, solar
 191 flares can modulate the electrodynamics of the ionosphere (Qian et al., 2012; Zhang
 192 et al., 2017). As seen in Figure 3, the JRO observations of vertical plasma drift velocity
 193 (blue) show a clear response to the solar flare, and the drifts exhibit a sudden ~ 15
 194 ms^{-1} decrease at the onset of the solar flare. The WACCM-X simulations only exhibit
 195 a weak ($1\text{-}2 \text{ms}^{-1}$) response to the solar flare at 285°E geographic longitude (black). How-
 196 ever, a stronger response occurs in the WACCM-X simulations at 320°E geographic lon-
 197 gitude (red), though it is still slightly weaker than seen in the observations. Nonethe-
 198 less, the vertical plasma drift response at 320°E is generally consistent with the JRO ob-
 199 servations, and we can thus use the simulations to understand the mechanism behind
 200 the rapid decrease in the vertical plasma drift during the solar flare.

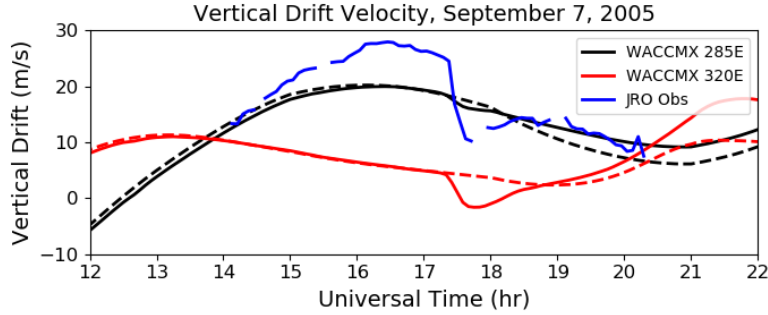


Figure 3. Equatorial vertical drift velocity on September 7, 2005 from Jicamarca 150-km echoes (blue), and WACCM-X simulations at 285°E (black) and 320°E (red) geographic longitude. Dashed lines indicate WACCM-X results without inclusion of the solar flare.

201 Previous studies investigating the solar flare effects on electrodynamics, and iono-
 202 spheric currents, have attributed the response to a change in the ionospheric conductiv-
 203 ity (Qian et al., 2012; Annadurai et al., 2018) and/or penetration electric field due to
 204 the imbalance of high latitude region-1 and region-2 field aligned currents. The later mech-
 205 anism was proposed by Zhang et al. (2017) as a source of the decrease in vertical plasma
 206 drift observed during the September 7, 2005 solar flare. The WACCM-X simulation does
 207 not include the effects of penetration electric fields, and we therefore attribute the change
 208 in vertical plasma drifts to changes in the ionospheric conductivity. It should be noted
 209 that we cannot entirely discount effects of penetration electric fields, and inclusion of pen-
 210 etration electric fields could lead to a larger vertical plasma drift response. The fact that
 211 the WACCM-X simulations capture a decrease in vertical plasma drifts at 320°E does,
 212 however, indicate that conductivity changes are an important mechanism by which solar
 213 flares influence electrodynamics.

214 The changes in the WACCM-X Hall (σ_H) and Pedersen (σ_P) conductivities at 17:45
 215 UT are shown in Figure 4 for 285°E and 320°E geographic longitude. Note that the changes
 216 are calculated relative to a WACCM-X simulation that did not include the solar flare
 217 forcing. For reference, maximum Hall and Pedersen conductivities at this time in the WACCM-
 218 X simulation without the solar flare are $\sim 8 \times 10^{-4}$ S/m and $\sim 5 \times 10^{-4}$ S/m, respectively.
 219 The conductivity changes due to the solar flare are thus large compared to the background
 220 conductivities. The corresponding zonal winds are shown in Figures 4c and 4f. Note that
 221 the zonal winds are largely unchanged by the solar flare below ~ 175 km, and are enhanced
 222 by $5\text{-}10 \text{ ms}^{-1}$ above 200 km (not shown). The change in Hall conductivity due to the
 223 flare is larger at 285°E than it is at 320°E, which should contribute to a larger decrease
 224 in the daytime eastward electric field, and thus a larger decrease in the vertical drift at
 225 285°E. The change in Pedersen conductivity due to the solar flare is generally similar
 226 at the two longitudes. The background zonal winds are, however, notably different which
 227 is likely due to the differences in local time at the two longitudes (12:45 SLT at 285°E
 228 and 15:05 SLT at 320°E). We therefore attribute the smaller change in the simulated drift
 229 response at 285°E to be due to the zonal winds at the time of the solar flare, and it is
 230 possible that WACCM-X does not capture the flare effects at 285°E due to deficiencies
 231 in the zonal winds. These differences highlight the need to accurately simulate both the
 232 neutral winds and conductivities in order to accurately simulate the solar flare effects
 233 on ionospheric electrodynamics.

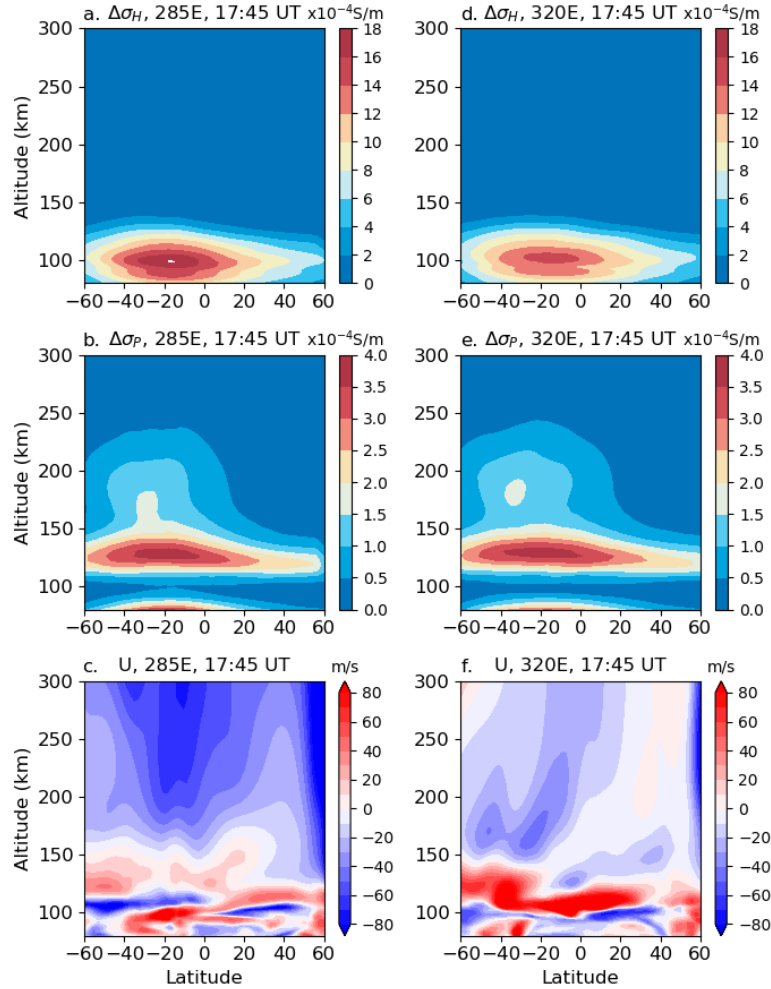


Figure 4. Changes in (a) Hall and (b) Pedersen conductivity at 285°E geographic longitude and 17:45 UT. (c) Zonal wind at 285° geographic longitude and 17:45 UT. (d-f) Same as (a-c) except for at 320°E geographic longitude.

5 Conclusions

The present study investigates the effects of the September 7, 2005 X-17 solar flare on the equatorial ionosphere using a combination of JRO observations and WACCM-X simulations. The solar flare is found to produce similar changes in the layering structure of observed 150-km echoes and simulated electron densities. In particular, both reveal a rapid descent at the onset of the solar flare, followed by a gradual ascent following the solar flare. The 150-km echo layers and contours of constant electron density are also both found to be narrower in vertical extent following the solar flare. These similarities support a connection between the background electron density and the layering structure that is seen in 150-km echoes. The reason for this relationship does, however, remain unknown, and further investigations into this connection will help in understanding the mechanisms that form the still unexplained 150-km echoes. The results also demonstrate that relatively coarse resolution whole atmosphere-ionosphere general circulation models, such as WACCM-X, can provide insight into smaller-scale structures in the equatorial ionosphere. This represents a new application of such models, enabling

249 potential future investigations focused on understanding, for example, the day-to-day
250 variability of 150-km echoes.

251 The effect of the solar flare on the equatorial vertical plasma drifts was also inves-
252 tigated. The JRO observations show a sudden decrease in vertical plasma drift veloc-
253 ity of 15-20 ms⁻¹ after the onset of the solar flare. The WACCM-X simulations repro-
254 duce a decrease in vertical plasma drift at 320°E geographic longitude, but only a weak
255 (1-2 ms⁻¹) decrease at the longitude of Jicamarca (285°E). The vertical plasma drift changes
256 are **primarily** attributed to changes in the conductivity in the simulations, which changes
257 the daytime eastward electric field, and the longitudinal differences may be related to
258 differences in the zonal winds at the time of the solar flare. This demonstrates that sim-
259 ulating the electrodynamic effects of solar flares requires accurately simulating both the
260 zonal winds as well as the conductivities. **Penetration electric fields may also in-**
261 **fluence the response of the equatorial vertical plasma drifts to the solar flare,**
262 **though the present results suggest that this may be a secondary effect.**

263 Acknowledgments

264 WACCM-X is part of the Community Earth System Model (CESM) and the source code
265 is available at <http://www.cesm.ucar.edu>. The WACCM-X simulation output, and Ji-
266 camarca SNR observations, used in this publication are available at <https://doi.org/10.26024/ahcm-6d40>. The GOES x-ray observations are available from NASA NCEI
267 (<https://www.ngdc.noaa.gov/stp/satellite/goes/dataaccess.html>). We would
268 like to acknowledge high-performance computing support from Cheyenne (doi:10.5065/D6RX99HX)
269 provided by NCAR's Computational and Information Systems Laboratory. This materi-
270 al is based upon work supported by the National Center for Atmospheric Research, which
271 is a major facility sponsored by the National Science Foundation under Cooperative Agree-
272 ment No. 1852977. We thank the International Space Science Institute for facilitating
273 discussions related to this paper as part of the International Team "An Exploration of
274 the Valley Region in the Low Latitude Ionosphere: Response to Forcing from Below and
275 Above and Relevance to Space Weather". The participation of J. L. C. in this work is
276 part of the project supported by the Deutsche Forschungsgemeinschaft (DFG, German
277 Research Foundation) under SPP 1788 (DynamicEarth)-CH 1482/1-2 (DYNAMITE2).
278

279 References

- 280 Annadurai, N. M. N., Hamid, N. S. A., Yamazaki, Y., & Yoshikawa, A. (2018).
281 Investigation of Unusual Solar Flare Effect on the Global Ionospheric Cur-
282 rent System. *J. Geophys. Res. Sp. Phys.*, *123*(10), 8599–8609. doi:
283 10.1029/2018JA025601
- 284 Balsley, B. B. (1964). Evidence of a stratified echoing region at 150 kilometers in
285 the vicinity of the magnetic equator during daylight hours. *J. Geophys. Res.*,
286 *69*(9), 1925–1930. doi: 10.1029/JZ069i009p01925
- 287 Chamberlin, P. C., Woods, T. N., & Eparvier, F. G. (2008). Flare Irradiance Spec-
288 tral Model (FISM): Flare component algorithms and results. *Sp. Weather*,
289 *6*(5). doi: 10.1029/2007SW000372
- 290 Chau, J., Woodman, R., Milla, M., & Kudeki, E. (2009). Naturally enhanced ion-
291 line spectra around the equatorial 150-km region. *Ann. Geophys.*, *27*, 933–942.
- 292 Chau, J. L. (2004). Unexpected spectral characteristics of VHF radar signals from
293 150-km region over Jicamarca. , *31*, L23803, doi:10.1029/2004GL021620.
- 294 Chau, J. L., & Kudeki, E. (2006). Statistics of 150-km echoes over Jicamarca based
295 on low-power VHF observations. *Ann. Geophys.*, *24*(5), 1305–1310. doi: 10
296 .5194/angeo-24-1305-2006
- 297 Chau, J. L., & Kudeki, E. (2013). Discovery of two distinct types of equatorial
298 150 km radar echoes. *Geophysical Research Letters*, *40*(17), 4509–4514.
299 Retrieved from <https://agupubs.onlinelibrary.wiley.com/doi/abs/>

- 300 10.1002/grl.50893 doi: 10.1002/grl.50893
- 301 Chau, J. L., Kudeki, E., & Milla, M. (2009). Multi-frequency radar studies of the
302 equatorial 150-km region. In N. Swarnalingam & W. K. Hocking (Eds.), *Pro-*
303 *ceedings of the 12th international symposium on equatorial aeronomy* (p. 165-
304 168). London, Ontario, Canada.
- 305 Choudhary, R. K., St.-Maurice, J.-P., & Mahajan, K. K. (2004). Observa-
306 tion of coherent echoes with narrow spectra near 150 km altitude during
307 daytime away from the dip equator. *Geophys. Res. Lett.*, *31*(19). doi:
308 10.1029/2004GL020299
- 309 de Paula, E. R., & Hysell, D. L. (2004). The São Luís 30 MHz coherent scatter
310 ionospheric radar: System description and initial results. *Radio Sci.*, *39*(1).
311 Retrieved from [https://agupubs.onlinelibrary.wiley.com/doi/abs/](https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2003RS002914)
312 [10.1029/2003RS002914](https://doi.org/10.1029/2003RS002914) doi: 10.1029/2003RS002914
- 313 Gelaro, R., McCarty, W., Suárez, M. J., Todling, R., Molod, A., Takacs, L., ...
314 Zhao, B. (2017). The Modern-Era Retrospective Analysis for Research and
315 Applications, Version 2 (MERRA-2). *J. Clim.*, *30*(14), 5419–5454. doi:
316 10.1175/JCLI-D-16-0758.1
- 317 Heelis, R. A., Lowell, J. K., & Spiro, R. W. (1982). A model of the high-latitude
318 ionospheric convection pattern. *J. Geophys. Res. Sp. Phys.*, *87*(A8), 6339–
319 6345. doi: 10.1029/JA087iA08p06339
- 320 Kudeki, E., & Fawcett, C. D. (1993). High resolution observations of 150
321 km echoes at Jicamarca. *Geophys. Res. Lett.*, *20*(18), 1987–1990. doi:
322 10.1029/93GL01256
- 323 Kudeki, E., Fawcett, C. D., Ecklund, W. L., Johnston, P. E., & Franke, S. J. (1998).
324 Equatorial 150-km irregularities observed at Pohnpei. *Geophys. Res. Lett.*,
325 *25*(21), 4079–4082. doi: 10.1029/1998GL900069
- 326 Lehmacher, G., Lu, X., Kudeki, E., Reyes, P. M., & Milla, M. (2018, May). Si-
327 multaneous observations of 150-km echoes and ionosonde virtual heights at
328 jicamarca. In *Proceedings of the 12th international workshop on technical and*
329 *scientific aspects of mst radar*. Ahmedabad, India.
- 330 Lehmacher, G. A., Kudeki, E., Akgiray, A., Guo, L., Reyes, P., & Chau, J. (2009).
331 Radar cross sections for mesospheric echoes at jicamarca. *Annales Geophysi-*
332 *cae*, *27*(7), 2675–2684. Retrieved from [https://www.ann-geophys.net/27/](https://www.ann-geophys.net/27/2675/2009/)
333 [2675/2009/](https://doi.org/10.5194/angeo-27-2675-2009) doi: 10.5194/angeo-27-2675-2009
- 334 Lehmacher, G. A., Kudeki, E., Reyes, P. M., Lee, K., Heale, C. J., & Snively, J. B.
335 (2019). Gravity wave ducting observed in the mesosphere over jicamarca,
336 peru. *Journal of Geophysical Research: Atmospheres*, *124*(10), 5166–5177.
337 Retrieved from [https://agupubs.onlinelibrary.wiley.com/doi/abs/](https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2019JD030264)
338 [10.1029/2019JD030264](https://doi.org/10.1029/2019JD030264) doi: 10.1029/2019JD030264
- 339 Liu, H.-L., Bardeen, C. G., Foster, B. T., Lauritzen, P., Liu, J., Lu, G., ... Wang,
340 W. (2018). Development and Validation of the Whole Atmosphere Community
341 Climate Model With Thermosphere and Ionosphere Extension (WACCM-X
342 2.0). *J. Adv. Model. Earth Syst.*, *10*(2), 381–402. doi: 10.1002/2017MS001232
- 343 Liu, J., Liu, H., Wang, W., Burns, A. G., Wu, Q., Gan, Q., ... Schreiner, W. S.
344 (2018). First Results From the Ionospheric Extension of WACCM-X During
345 the Deep Solar Minimum Year of 2008. *J. Geophys. Res. Sp. Phys.*, *123*(2),
346 1534–1553. doi: 10.1002/2017JA025010
- 347 Marsh, D. R., Mills, M. J., Kinnison, D. E., Lamarque, J. F., Calvo, N., &
348 Polvani, L. M. (2013). Climate change from 1850 to 2005 simulated in
349 CESM1(WACCM). *J. Clim.*.. doi: 10.1175/JCLI-D-12-00558.1
- 350 Neale, R. B., Richter, J., Park, S., Lauritzen, P. H., Vavrus, S. J., Rasch, P. J.,
351 & Zhang, M. (2013). The Mean Climate of the Community Atmosphere
352 Model (CAM4) in Forced SST and Fully Coupled Experiments. *J. Clim.*.. doi:
353 10.1175/JCLI-D-12-00236.1
- 354 Oppenheim, M. M., & Dimant, Y. S. (2016). Photoelectron-induced waves: A likely

- 355 source of 150 km radar echoes and enhanced electron modes. *Geophys. Res.*
 356 *Lett.*, *43*(8), 3637–3644. doi: 10.1002/2016GL068179
- 357 Patra, A., Pavan Chaitanya, P., St-Maurice, J.-P., Otsuka, Y., Yokoyama, T., & Ya-
 358 mamoto, M. (2017). The solar flux dependence of ionospheric 150 km radar
 359 echoes and implications. *Geophysical Research Letters*, *44*(22).
- 360 Patra, A. K., Pavan Chaitanya, P., & Tiwari, D. (2011). Characteristics of 150
 361 km echoes linked with solar eclipse and their implications to the echoing phe-
 362 nomenon. *J. Geophys. Res. Sp. Phys.*, *116*(A5). doi: 10.1029/2010JA016258
- 363 Patra, A. K., Yokoyama, T., Otsuka, Y., & Yamamoto, M. (2008). Daytime 150-km
 364 echoes observed with the Equatorial Atmosphere Radar in Indonesia: First
 365 results. *Geophys. Res. Lett.*, *35*(6). doi: 10.1029/2007GL033130
- 366 Perkins, F., Salpeter, E., & Yngvesson, K. (1965). Incoherent scatter from plasma
 367 oscillations in the ionosphere. *Physical Review Letters*, *14*(15), 579.
- 368 Pettit, J., Randall, C. E., Marsh, D. R., Bardeen, C. G., Qian, L., Jackman, C. H.,
 369 ... Harvey, V. L. (2018). Effects of the September 2005 Solar Flares and Solar
 370 Proton Events on the Middle Atmosphere in WACCM. *J. Geophys. Res. Sp.*
 371 *Phys.*. doi: 10.1029/2018JA025294
- 372 Qian, L., Burns, A. G., Chamberlin, P. C., & Solomon, S. C. (2011). Variability
 373 of thermosphere and ionosphere responses to solar flares. *J. Geophys. Res. Sp.*
 374 *Phys.*, *116*(A10). doi: 10.1029/2011JA016777
- 375 Qian, L., Burns, A. G., Solomon, S. C., & Chamberlin, P. C. (2012). Solar flare im-
 376 pacts on ionospheric electrodynamics. *Geophys. Res. Lett.*, *39*(6). doi: 10.1029/
 377 2012GL051102
- 378 Reyes, P. (2012). *Solar Flare Effects Observed Over Jicamarca During MST-ISR*
 379 *Experiments* (Master of Science, University of Illinois at Urbana-Champaign).
 380 Retrieved from <http://hdl.handle.net/2142/31196>
- 381 Reyes, P. (2017). *Study of Waves Observed in the Equatorial Ionospheric Val-*
 382 *ley Region using Jicamarca ISR and VIPIR Ionosonde* (Doctor of Phi-
 383 losophy, University of Illinois at Urbana-Champaign). Retrieved from
 384 <http://hdl.handle.net/2142/98349>
- 385 Richmond, A. D., Ridley, E. C., & Roble, R. G. (1992). A thermosphere/ionosphere
 386 general circulation model with coupled electrodynamics. *Geophys. Res. Lett.*,
 387 *19*(6), 601–604. doi: 10.1029/92GL00401
- 388 Roble, R. G., Ridley, E. C., Richmond, A. D., & Dickinson, R. E. (1988). A cou-
 389 pled thermosphere/ionosphere general circulation model. *Geophys. Res. Lett.*,
 390 *15*(12), 1325–1328. doi: 10.1029/GL015i012p01325
- 391 Smith, A. K., Pedatella, N. M., Marsh, D. R., & Matsuo, T. (2017). On the Dynam-
 392 ical Control of the Mesosphere–Lower Thermosphere by the Lower and Middle
 393 Atmosphere. *J. Atmos. Sci.*, *74*(3), 933–947. doi: 10.1175/JAS-D-16-0226.1
- 394 Tsunoda, R. T., & Ecklund, W. L. (2008). On the sheet-like nature of 150 km (F1)
 395 radar echoes. *Geophys. Res. Lett.*, *35*(5). doi: 10.1029/2007GL032152
- 396 Zhang, R., Liu, L., Le, H., & Chen, Y. (2017). Equatorial ionospheric electrodynami-
 397 cally during solar flares. *Geophys. Res. Lett.*, *44*(10), 4558–4565. doi: 10.1002/
 398 2017GL073238