

1 **Arctic sea-ice loss fuels extreme European snowfall**

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24 **The accelerated loss of Arctic sea-ice has been implicated with severe cold and snowy**
25 **mid-latitude winters. However, the mechanisms and a direct link remain elusive due to**
26 **limited observational evidence. Here we present atmospheric water vapour isotope**
27 **measurements from Arctic Finland during “the Beast from the East” - a severe anticyclonic**
28 **outbreak that brought heavy snowfall and freezing across Europe in February 2018. We**
29 **find that an anomalously warm Barents Sea, with a 60% ice-free surface, supplied up to**
30 **9.3 mm d⁻¹ moisture flux to this cold north-easterly airflow. We demonstrate that**
31 **approximately 140 gigatonnes of water was evaporated from the Barents Sea during the**
32 **event, supplying up to 88% of the corresponding fresh snow over Northern Europe.**
33 **Reanalysis data show that from 1979 to 2020, net March evaporation across the Barents**
34 **Sea increased by approximately 70 kg per square metre of sea-ice lost ($r^2=0.73$, $p<0.01$),**
35 **concurrent with a 1.6 mm (water equivalent) per year increase in Europe’s maximum**
36 **snowfall. Our analysis directly links Arctic sea-ice loss with increased evaporation and**
37 **extreme snowfall, and signifies that by 2080, an *Atlantified* ice-free Barents Sea will be a**
38 **major source of winter moisture for continental Europe.**

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40 Arctic sea-ice plays a critical role in the hydrological cycle and global climate system¹. In
41 particular, the areal extent and concentration of sea ice controls thermodynamic and
42 radiative processes driving water vapour, clouds, and aerosol feedbacks². Accelerated sea
43 ice loss over recent decades³ has been linked to increased surface evaporation and latent
44 heat flux^{4,5}, as well as localised increases in cloud formation, precipitation, and radiation
45 absorption that further amplifies Arctic warming⁶⁻⁸.

46

47 The Barents Sea is an Arctic hotspot⁹ where maximum (March) winter sea ice cover
48 has decreased by 54% since 1979¹⁰ – an area of ~570,000 km². This sea ice decline of 11,200
49 km² per year³ has been accompanied by increasing snow mass trends across large areas of
50 Eurasia that are adjacent to the Arctic Ocean¹¹, particularly in autumn^{11–13}. Evidence further
51 suggests a dynamic link whereby autumn Barents-Kara sea-ice and snow cover anomalies in
52 Eurasia can force extreme cold and snowy mid-latitude winters^{12,14–18}. In part, this lagged
53 connection reflects the vertical propagation of surface energy to the stratosphere, which
54 can weaken the winter stratospheric polar vortex and induce strong anticyclonic flow over
55 the Arctic Ocean^{17,18}. These circulation anomalies manifest as phases of negative Arctic
56 Oscillation (AO–) and North Atlantic Oscillation (NAO–)^{19,20}, that can drive cold air advection
57 and heavy snowfall across continental mid-latitudes, such as in winters 2009-10, 2010-11,
58 2012-13 and 2017-18^{21–24}. However, a direct link between winter Barents sea-ice loss and
59 the recent extreme snowy European winters²⁵ has yet to be substantiated.

60

61 Here, we present empirical evidence for the direct impact of Arctic sea-ice decline on
62 a severe winter weather event in Europe during February-March 2018. Popularly dubbed
63 “*the Beast from the East*”, the event coincided with strong NAO– circulation and drove
64 anomalous low surface air temperatures (SAT) and heavy snowfall across Europe with
65 severe socio-economic impacts^{22,23} (Fig. 1). We captured this event with continuous *in situ*
66 measurements of atmospheric water vapour isotopes ($\delta^{18}\text{O}$, $\delta^2\text{H}$) in Arctic Finland, providing
67 unique observational constraints on the sea ice, moisture source, and kinematic processes
68 during the event^{26–28}. Specifically, we test the hypothesis that winter Barents sea-ice decline
69 – and the expansion of warm open water – provides an incipient evaporative source of
70 atmospheric moisture that can directly fuel heavy snowfalls over Europe.

71

72 **The Beast from the East**

73 The spatial development of weekly composited sea level pressure (SLP) and SAT anomalies
74 encompassing the February-March 2018 event are presented in Supplementary Figure 1. A
75 major atmospheric precursor was the displacement and subsequent split of the
76 stratospheric polar vortex on 11 February 2018, coincident with a major Sudden
77 Stratospheric Warming (SSW) event^{22,29}. The polar vortex split induced a dynamic reversal of
78 the westerly zonal-mean winds poleward of 60°N at 10 hPa, triggering a marked
79 temperature increase (+33 °C) over the polar stratosphere lasting ~20 days²². The successive
80 downward migration of easterly winds to the troposphere²² favoured the development of a
81 strong NAO– surface response with a large region of high pressure over Northern
82 Scandinavia and the Barents Sea²³ (Fig. 1a). Between 19 February and 5 March this strong
83 anticyclone steered Arctic airflow directly into Europe, driving extreme negative SAT
84 anomalies and heavy snowfall (Fig. 1). Blizzards were acutely disruptive across the British
85 Isles and Western Europe where the Arctic outbreak converged with a deep Atlantic cyclone
86 in early March (“Storm Emma”)²³. Europe was subsequently hit by a second cold wave
87 beginning 14 March when a consecutive anticyclone anchored over Scandinavia and drove
88 Arctic airflow across the continent²⁹ (Supplementary Figure 1).

89

90 Stable isotope and automated weather station (AWS) measurements were recorded
91 at the Finnish Meteorological Institute’s Sammaltunturi station in Pallas-Yllästunturi
92 National Park, Arctic Finland (hereafter “Pallas”; 67.973 °N, 24.116 °E; 565 m asl), and
93 capture the 7-week period before and after the polar vortex split (see Methods). AWS
94 measurements indicate the build-up of high-pressure beginning 11 February coincident with

95 the SSW and anticyclogenesis over the Barents Region (Fig. 2 and Supplementary Fig. 1).
96 Maximum barometric pressure at Pallas was attained on 28 February (980 hPa), equivalent
97 to 1055 hPa SLP and coincident with the NAO– minima on 1 March (-1.7σ) (Fig. 2).

98

99 Prior to the event, mean vapour $\delta^{18}\text{O}$, $\delta^2\text{H}$, and deuterium (d)-excess values at Pallas
100 were $-28.6 \text{ ‰} (\pm 4.3)$, $-214.6 \text{ ‰} (\pm 36.3)$ and $14.1 \text{ ‰} (\pm 4.9)$ respectively (21 December 2017
101 – 18 February 2018) (Fig. 2), where d-excess is calculated as $\delta^2\text{H} - 8 \cdot \delta^{18}\text{O}^{30}$. While meteoric
102 δ -values reflect evaporation and condensation processes related to air mass temperature
103 and saturation levels^{30,31}; d-excess reflects and retains information pertaining to conditions
104 at the evaporative source, primarily relative humidity (RH) and sea surface temperature
105 (SST)^{32,33}. Accordingly, Pallas $\delta^{18}\text{O}$ values were positively correlated with local SAT ($r^2=0.5$,
106 $p<0.05$). Moreover, whilst the lowest measured d-excess value (1.3 ‰) coincided with the
107 highest local SAT on 11 January (1.0°C), maximum d-excess did not correspond to the lowest
108 SAT on 24 January (-24.4°C). Instead, beginning 19 February we captured a remarkable 16.0
109 ‰ abrupt increase in d-excess at Pallas, culminating in maximum d-excess on 23 March
110 (31.2‰), and coinciding with increased snowfall over Northern Europe ($15\text{-}60^\circ\text{E}$, $50\text{-}70^\circ\text{N}$;
111 Fig. 2c) and onset of the Beast from the East.

112

113 **Moisture flux at the sea-ice margin**

114 We used Lagrangian atmospheric back-trajectory modelling driven by the Global Data
115 Assimilation System³⁴ to diagnose the Pallas vapour origin and transport processes during
116 the event (see Methods). Over the 14-week measurement period, our model output
117 indicates four evaporative source regions: Barents (32 %), Atlantic (27 %), Eurasia (22 %) and
118 Baltic (19 %) (Supplementary Fig. 2). However, northerly trajectories and snowfall during the

119 event were supplied by moisture originating from the Barents Sea (Fig. 3a). At this time, the
120 anticyclone drove cold, dry air over the ice pack across open water, where our modelled
121 trajectories demonstrate a sharp increase in specific humidity due to intense evaporation
122 along the sea ice edge ($+0.5 \text{ g kg}^{-1}$ per air parcel) (Fig. 3a). There is a corresponding
123 deepening of the atmospheric boundary layer (ABL) from $\sim 100 \text{ m}$ over the sea ice, to 2000
124 m-thick downstream of the ice margin due to strong turbulent and radiative processes
125 (Supplementary Fig. 3). These processes are captured in MODIS satellite imagery as parallel
126 cloud bands (or “cloud streets”) that formed off the ice margin (Fig. 3b), indicating the
127 development of a strongly heated convective ABL with intense surface evaporation and
128 condensation. Low level airflow advected the moisture south across the north Norwegian
129 coast and inland to Pallas (Fig. 3c), with a total transport time from Barents moisture source
130 to sink of < 72 hours (minimum 28 hr) (Supplementary Fig. 3).

131

132 Using our back-trajectory output we manually identify three discrete pulses of
133 Barents Sea moisture advected to Europe during this period: 19 February–4 March, 14–20
134 March, and 23–28 March, each corresponding to increased continental snowfall and high
135 vapour and precipitation d-excess (Fig. 2b,c and Supplementary Fig. 4). Meteoric vapour
136 with a high d-excess reflects the slower diffusivity of the H_2^{18}O molecule during rapid
137 evaporation, when there is insufficient time for vapour to reach isotopic equilibrium with
138 the ocean surface³³. We hypothesise that such conditions, driven by large RH gradients and
139 strong surface winds, dominated the Barents sea-ice margin to drive rapid high d-excess
140 vapour production^{27,31}. Critically, as the polar air mass was close to saturation over the ice
141 pack (98% RH), the anomalously warm Barents SSTs ($2\text{-}5 \text{ }^\circ\text{C}$) (Fig. 1b) coupled with low near-
142 surface RH (65-70%) induced rapid kinetic (non-equilibrium) isotope fractionation as the

143 northerly air flowed across the sea-ice/open-water boundary^{27,31}. Furthermore, compared
144 to atmospheric vapour originating from the Baltic and Eurasia regions, we find that Barents
145 evaporate is characterised by relatively high mean $\delta^{18}\text{O}$ values that are comparable to the
146 Atlantic-derived evaporate (Supplementary Fig. 2). We attribute this to the well-
147 documented increased inflow of Atlantic water into the Barents Sea - the primary driver for
148 marked sea ice loss in this Arctic warming hotspot⁹.

149

150 **Arctic sea-ice and European snowfall trends**

151 Our analyses identify the Barents Sea as a key source of atmospheric moisture during the
152 2018 event. Fundamental to this process was a ~60% ice-free, *Atlantified* sector of the
153 Barents that enabled the direct transfer of latent energy from the ocean surface to the
154 lower atmosphere. Using ERA5 reanalysis data³⁵, we determine a net Barents moisture flux
155 of ~140 gigatonnes (Gt) to the atmosphere during the “Beast from the East” (19 February–
156 28 March), attaining a maximum evaporation rate of 9.3 mm d⁻¹ on 1 March coincident with
157 the NAO minima (Fig. 2e). Global Snow Monitoring observations (GlobSnow)³⁶ show a net
158 snow mass increase of 159.2 ± 2.9 Gt (water equivalent) across Northern Europe over this
159 38-day period (Fig. 1b), indicating that Barents evaporation potentially contributed up to
160 ~88% to this fresh snow cover. By restricting our analysis to the initial pulse of Barents
161 moisture advected to Northern Europe between 19 February and 4 March (Fig. 2, “Pulse 1”)
162 – thereby excluding the Atlantic influence from “Storm Emma”²³ – we find that Barents
163 evaporation potentially supplied 54 Gt moisture to the atmosphere, equivalent to ~69% of
164 Northern Europe’s net snow increase of 78.8 ± 1.4 Gt³⁶ (Fig. 3e). This equates to a mean
165 evaporative flux of 3.9 Gt d⁻¹ from the Barents during Pulse 1 (Fig. 3c) that, under the
166 prevailing northerly airflow, contributed ~18.5% of the 21 Gt daily mean total column water

167 vapour (TCWV) over Northern Europe (surface to 850 hPa). By comparison, during a
168 northerly winter outbreak from 13–19 March 1979³⁷, when Barents sea-ice cover was
169 ~640,000 km² (56%) more extensive, a mean flux of only 1.4 Gt d⁻¹ was evaporated off the
170 Barents Sea (Fig. 3d). This represents ~4% of Northern Europe’s mean TCWV budget during
171 the 1979 event, and the corresponding snowfall accumulation was around half (41.5 ± 0.8
172 Gt water equivalent) of that compared with Pulse 1 of the “Beast from the East”.

173

174 Since 1979 the Barents Sea has been responsible for 95% of the observed March sea-
175 ice loss across the entire Arctic³⁸. We use satellite observations of sea-ice¹⁰ and ERA5
176 reanalysis³⁵ to investigate long-term dynamic links with atmospheric moistening in the
177 Barents Region^{2,6}, as well as increasing European extreme snowfall²⁵ (see Methods).
178 Between 1979-2020, we find a linear March Barents sea-ice decline of ~11,200 km² yr⁻¹ ($r^2 =$
179 0.66, $p < 0.05$) consistent with earlier estimates³, and a corresponding net March
180 evaporation increase of 1.01 Gt yr⁻¹ ($r^2 = 0.74$, $p < 0.05$) (Fig. 4a). Over this period, Barents
181 surface evaporation is negatively correlated with sea-ice area ($r^2 = 0.73$, $p < 0.01$),
182 demonstrating a mean net March evaporative flux increase of 69.9 Gt per 1 million km² sea-
183 ice loss (~70 kg per m²) (Fig. 4a). Moreover, we find that while mean March snowfall across
184 Northern Europe has decreased by 8.2 mm (water equivalent) per decade since 1979, the
185 maximum March snowfall – indicative of extreme heavy snowfall events²⁴ – has increased
186 by 16.0 mm per decade ($r^2 = 0.42$, $p < 0.05$), and is linearly congruent with increased Barents
187 evaporation ($r^2 = 0.52$, $p < 0.05$) (Fig. 4a and Supplementary Fig. 5).

188

189 Previous studies also reveal a strong connection between Arctic sea-ice loss and
190 increased localised evaporation in the Barents Region^{4,6}, yet the long-term relationship has

191 not been quantified across the Arctic, nor the direct and non-lagged seasonal link
192 established with European winter snowfall. We explore these relationships by spatially
193 regressing winter (December-March) ERA5 reanalysis fields³⁵ and sea ice observations¹⁰
194 from 1979 to 2020. All fields are detrended and only relationships significant at the 95%
195 confidence level considered (see Methods). Since 1979, winter surface evaporation across
196 the Barents Region is negatively correlated with winter sea-ice area, with strong
197 relationships also apparent in the Chukchi and Bering Seas ($r^2 > -0.5$; $p < 0.05$) (Fig. 4b).
198 Moreover, we find that European snowfall anomalies positively correlate with Barents
199 evaporation, whereby enhanced evaporation increases heavy snowfall across Northern
200 Europe (Fig. 4c). Whilst these relationships are particularly robust across Fennoscandia with
201 coefficients up to $r^2=0.7$ ($p < 0.05$), the influence of declining Barents sea-ice also spans a 20°
202 latitudinal range across Northern Europe, extending across the Baltic states and south into
203 Russia (Fig. 4c).

204

205 **Implications of projected ice-loss**

206 The Barents Sea was a hotbed of extreme moisture flux during the Beast from the East in
207 2018. The pulses of high d-excess moisture that we observe being advected from the
208 Barents into Europe represent a ‘smoking gun’ confirming our back-trajectory analyses that
209 the vapour originated from increasingly exposed Arctic waters. Given projections of a winter
210 ice-free Barents Sea by 2061-2088³⁹, our observations support a future increase of locally-
211 sourced high-latitude atmospheric moisture and precipitation across the Arctic^{4,40,41}. Our
212 analysis further signifies an increased potential for extreme winter snowfall across Northern
213 Europe where temperatures remain sufficiently cold to yield snow, for example under
214 prevailing cold northerly airflow conditions such as those associated with weakened

215 stratospheric polar vortex events and/or NAO– circulation^{13,17,42}. Whether or not this effect
216 is sustainable in the long-term under current and projected mean rates of warming^{1,43}
217 warrants further investigation, including the potential for moist Arctic air masses to
218 thermodynamically offset (or exceed) the atmospheric cooling associated with these
219 events^{43,44}. It is widely considered that the poleward transport and convergence of moisture
220 into the Arctic from lower latitudes is the fundamental mechanism for future Arctic
221 amplification of the hydrologic cycle^{45,46}. We conclude that an increasingly exposed, ice-free
222 Arctic Ocean also provides an important *local* supply of atmospheric moisture that is a major
223 source of winter precipitation for continental Europe.

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346 **Author Contributions**

347 H.B. conducted the research, created the figures, and wrote the manuscript. H.B., A.H.,
348 E.S.K. and J.M.W (Project PI) conceived and designed the study. J.M.W., E.S.K., K-R.M. and
349 H.M. conducted the fieldwork. K-R.M., E.S.K., H.M. and P.D.A. performed and/or
350 contributed to the isotope data measurements and post-processing. H.B. and A.H.
351 performed the back-trajectory and long-term analyses. All authors contributed comments
352 and/or revisions to the manuscript.

353

354 **Competing Interests**

355 The authors declare no competing interests.

356

357 **Figure captions**

358 **Figure 1. Synoptic climatology during the Beast from the East.** Maps show (a) daily
359 composited near surface (2m) air temperature (T) and sea level pressure anomalies, and (b)
360 sea surface temperature (SST) anomalies and total snowfall³⁶, over the period from 19
361 February to 28 March 2018. Anomalies are calculated relative to the 1981-2010 baseline³⁵.
362 High (H) and low (L) pressure centres are indicated in (a), and mean sea ice cover¹⁰ is
363 depicted in (b) as light grey shading. The black square indicates the Pallas field site location
364 in Arctic Finland.

365

366 **Figure 2. Observations during winter 2017-18.** Timeseries show Pallas (a) vapour (line) and
367 snow (circles) $\delta^{18}\text{O}$ (blue) and $\delta^2\text{H}$ (red), (b) vapour (blue lines) and snow (circles) d-excess,
368 (c) Pallas barometric pressure (570 m asl) and daily snowfall³⁶ (bars) over Northern Europe
369 (asterisks indicate missing data), (d) Pallas vapour mixing ratio (red) and air temperature
370 (blue), and (e) positive (red) and negative (blue) NAO index values. Pallas data represent 5-
371 minute averages, and daily mean vapour d-excess is also shown in (b) by the thick blue line.
372 Blue columns indicate three pulses of Barents moisture advected to Europe during the Beast
373 from the East (black dashed lines). NAO data in (e) were obtained from the National
374 Weather Service Climate Prediction Center (<https://www.cpc.ncep.noaa.gov/data>).

375

376 **Figure 3. Barents Sea moisture advection to Northern Europe.** (a) Back-trajectories from
377 Pallas (square) between 19-27 February 2018 and associated mean vapour d-excess and
378 $\delta^{18}\text{O}$. Colours depict hourly specific humidity changes (Δq), where a positive (negative) Δq
379 indicates a moisture increase (decrease) due to evaporation (precipitation). Grey circles
380 indicate either no net moisture change or a change above the ABL; (b) Aqua-MODIS satellite
381 image showing Barents “cloud streets”; (c) daily mean evaporation during outbreaks in 2018
382 and (d) 1979, and (e) Northern Europe snow mass³⁶ increase during the interval in (d).
383 Black/blue solid lines in (a,c-e) show mean sea-ice edge¹⁰, grey isolines in c-e represent
384 mean SLP³⁵ (4 hPa intervals).

385

386 **Figure 4. Historical Arctic sea-ice and atmospheric moisture links.** (a) Linear trends and
387 regression (inset) of the 3-yr running mean March Barents sea-ice area (blue) and net
388 surface evaporation (red), and Northern Europe’s maximum March snow mass (grey bars)
389 between 1979-2020. Spatial regression coefficients of detrended winter (December-March)







